CLIMATE SCIENCE A perspective for business leaders



Statkraft | DNV GL | Shell | TCS | Xyntéo

CONTENTS

CLIMATE SCIENCE

A perspective for business leaders

Fore	eword	3	
1.	Executive summary	4	
2.	How the Earth is warming	6	
3.	The role of natural factors		
4.	The role of human activity	13	
5.	What the future holds 5a. Greenhouse gas emissions 5b. Warming 5c. Snow, ice and frozen ground	16 16 18 20	
6.	 Hazards 6a. Sea-level rise 6b. Heatwaves and drought 6c. Storms and floods 6d. Extreme waves 6e. Ocean acidification 6f. Food and water 6g. Tipping points 	24 26 27 28 29 31 33	
7.	The Fifth Assessment Report of the IPCC – The Physical Science Basis	36	
8.	Frequently asked questions	39	
Appendix: IPCC AR5 likelihood terminology			
Endi	notes	43	

Xyntéo alone is responsible for this document and any errors it contains.

© Copyright Xyntéo November 2013

Registered address: 3 Wesley Gate Queen's Road Reading RG1 4AP, UK

Registered in England Company number 5314641; VAT registration number 857 5824 79.

Designed and typeset by Soapbox, www.soapbox.co.uk Cover image: © WIN-Initiative/Getty Images 'Reflection on sea' Other images: www.istockphoto.com

Acknowledgements

CICERO played a key role in validating the climate science presented in this report. In particular we would like to acknowledge input from: Robbie Andrews, Christian Bjørnæs, Cecilie Mauritzen, Robert van Oort, Glen Peters, Pål Prestrud, Tiina Ruohonen, Bjørn H Samset and Maria Sand.

Xyntéo contributors: Stephen Battersby, Steve Esau, Saya Snow Kitasei and Gabrielle Walker.

SPONSORS

Istatkraft

Statkraft is Europe's largest generator of renewable energy and the leading power company in Norway. The company owns, produces and develops hydropower, wind power, gas-fired power and district heating. Statkraft is a major player in European power trading and has 3,600 employees in more than 20 countries. www.statkraft.com



As of September 2013, DNV and GL merged to form DNV GL. Driven by its purpose of safeguarding life, property and the environment, DNV GL enables organisations to advance the safety and sustainability of their business. The company provides classification and technical assurance along with software and independent expert advisory services to the maritime, oil and gas, and energy industries. It also provides certification services to customers across a wide range of industries. Operating in more than 100 countries, DNV GL's 16,000 professionals are dedicated to helping its customers make the world safer, smarter and greener www.dnv.com



Royal Dutch Shell is a global group of energy and petrochemicals companies with around 90,000 employees in more than 80 countries and territories. Shell's Upstream businesses explore for and extract crude oil and natural gas, often in joint ventures with international and national oil and gas companies. Shell's Downstream businesses manufacture, supply and market oil products and chemicals worldwide. Shell has been reporting on its environmental and social performance since 1997, and its sustainability performance is ranked in some leading indices.

www.shell.com

TATA CONSULTANCY SERVICES

Tata Consultancy Services (TCS) is an IT services, consulting and business solutions organisation that delivers real results to global business, ensuring a level of certainty no other firm can match. TCS offers a consulting-led, integrated portfolio of IT and IT-enabled infrastructure, engineering and assurance services. This is delivered through its unique Global Network Delivery Model, recognised as the benchmark of excellence in software development. A part of the Tata Group – India's largest industrial conglomerate – TCS has a global footprint and is listed on the National Stock Exchange and Bombay Stock Exchange in India. www.tcs.com

🔀 xyntéo

Xyntéo's mission is to 'reinvent growth'. As an advisory firm, it works with global companies to identify and carry out projects that aim to enable businesses to grow in a new way, fit for the resource, climate and demographic realities of the 21st century. The founder and engine of the GLTE partnership, Xynteo is also the driving force behind The Performance Theatre. The theatre is an annual meeting for CEOs and chairmen that aims to inspire the leadership needed to build a new kind of growth, capable of creating longer-term value for both shareholders and society as a whole. www.xynteo.com

GLTE

The GLTE partnership connects global businesses engaged in the pursuit of resource-efficient, low-carbon growth. It builds senior executives' knowledge of how the changing resource picture and climate change are affecting their businesses and industries, and of how other organisations are responding with innovation that embeds resource and carbon efficiency into the core of their businesses. What sets the GLTE partnership apart is its bias for action. The partnership conceives and conducts projects that aim to enable businesses to grow in a new way, fit for the resource, climate and demographic realities of the 21st century. Advisory firm Xyntéo founded and runs the GLTE partnership. www.xynteo.com/glte

FOREWORD



Climate change is rewriting the competitive landscape. Regulation, investment flows, and customer preferences – all are pushing businesses to reinvent the way they grow. Far-sighted business leaders are responding now to these forces. Not only do they see the risks that this transition will bring; more importantly, they also appreciate the opportunities that will be generated in the shift to a low-carbon economy.

However, many feel hampered by a widespread lack of understanding, or in some cases a deep scepticism relating to the underlying climate science that is driving this agenda. That is why I initiated this project, with the aim to provide an independent, clear, logical and compelling analysis of the state and frontiers of climate science, in a form that is tailored to the needs of business leaders, senior executives and other key stakeholders.

I am delighted that Royal Dutch Shell, DNV GL and Tata Consultancy Services have joined Statkraft in sponsoring this project. Together we hope that the information this document provides will help businesses create a hard leverage point for positive change based on the latest scientific frontiers, and enabling executives to make more informed decisions.

Christian Rynning-Tønnesen, CEO, Statkraft

This report will be periodically updated to reflect the latest climate science. Visit **www.xynteo.com** to read the up-to-date version.

1. EXECUTIVE SUMMARY

We need greenhouse gases. Carbon dioxide, water vapour and other trace gases act to trap heat and keep the planet warm. Without them Earth would be frozen and lifeless.

But you can have too much of a good thing. Since the Industrial Revolution, humans have been emitting large quantities of greenhouse gases, mainly by burning fossil fuels and turning forests into farmland. This has raised global temperature and had a host of other effects on our climate.

Despite its importance to informed decision-making, climate science has become a controversial and poorly communicated topic. Virtually all scientists agree that the climate is changing and that human activities are driving most of that change, while outside the scientific community these conclusions are often doubted. Most of the non-technical literature on climate change is often skewed by politics or prejudice.

We believe that business leaders need an unbiased view of the evidence, so our objective is to present the latest climate science in an accessible form and unfiltered through a political agenda.



The deadly and destructive Hurricane Sandy hit in 2012, becoming the second-costliest storm in US history.

> Climate change is already evident in many of Earth's systems. Global temperatures are rising, and so are sea levels. Glaciers are melting and permafrost is thawing. Oceans are becoming more acidic. Rainfall patterns are shifting, while some forms of extreme weather are becoming more frequent and severe.

These effects are almost certain to intensify over the 21st century. For businesses they will present new risks and new opportunities. Understanding the up-to-date scientific findings is critical for CEOs, senior executives and anyone who wants to remain informed and competitive in a warming world.

Large-scale reviews of climate science are published elsewhere, notably by the

Intergovernmental Panel on Climate Change (IPCC). The first part of their Fifth Assessment Report on climate science (AR5) was published in September 2013. This document is not intended to be an abbreviated version of that thousandpage review; rather it is a focused analysis of the latest findings that are most relevant for business. In assembling these findings, we have used scientific reports from both within and outside the IPCC, seeking out the most relevant and reliable research from many of the world's most respected scientific institutions. Nevertheless the IPCC's latest report is a convenient reference point and will be cited throughout the text.

KEY MESSAGES:

- Greenhouse gases emitted by human activities are responsible for most of the warming observed in recent decades. Although natural factors such as changes to the sun's output and volcanic activity also affect the climate, they cannot explain recent warming.
- Earth has warmed by about 0.85°C since the beginning of the last century. If carbon emissions follow the present high trajectory, by the end of this century temperatures are projected to reach 3.2 to 5.4°C above pre-industrial levels.
- Continued warming will have other effects on the climate system by the end of the 21st century, including:
 - A rise in global sea level that could reach 1.5 metres.
 - More heatwaves and droughts.
 - More storms and floods.
 - Impaired supplies of fresh water.
 - An increase in the ocean's acidity by 100 to 200% above pre-industrial levels.
 - Loss of ice, snow and permafrost.

— FAST FACT — The effects of climate change are almost certain to intensify over the 21st century

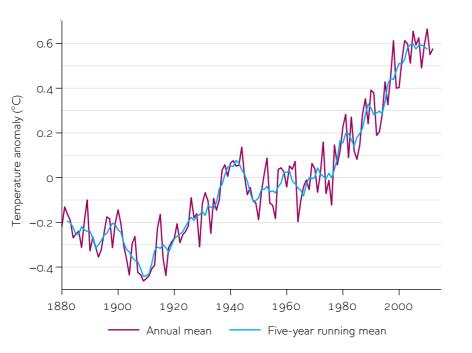
2. HOW THE EARTH IS WARMING

FIGURE 2: LOCAL WARMING 1901-2012

Overview

Average global temperature increased by about 0.85°C between 1901 and 2012. Most of that, around 0.72°C, has occurred since 1951 (see figure 1).1

FIGURE 1: GLOBAL TEMPERATURE



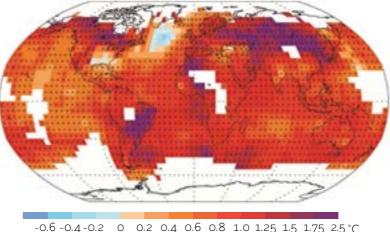
Temperatures measured by weather stations, ships and satellites, are used to calculate this global average across land and sea.

(Source: Adapted from NASA Goddard Institute for Space Studies. Plotted relative to the average temperature from 1951 to 1980.)

> Warming is not uniform (see figure 2). For example, Arctic temperatures have increased much faster than the global average rate over the past century.²

The rate of warming varies over time owing to short-term factors, including natural climate cycles.³ For example, there was little if any rise in average air temperatures in the period from 1998 to 2012,⁴ which is most likely the consequence of a cool phase in the tropical Pacific ocean (see section 3, 'The role of natural factors').

Oceans absorb heat and so act as a buffer for rising temperatures. Since 1955, the top 700 metres of the ocean has warmed by an average of 0.18°C and the top two kilometres by 0.09°C,⁵ accounting for 90% of the energy the Earth has gained since the 1960s.⁶ Heat is also reaching the deep ocean below 4,000 metres.⁷

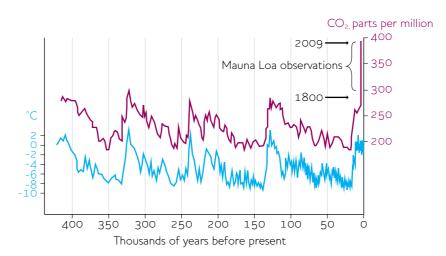


To put the rate of this recent warming in perspective, scientists have reconstructed past temperatures on many different timescales. Measurements from weather stations give a robust global temperature record back to about 1850. Further back in time scientists can calculate temperatures approximately using ice cores, ocean sediments, corals, tree rings and pollen buried in ancient lake mud.

Sediments hold the fossil skeletons of small organisms that preserve chemical clues to the climate going back many millions of years. They show that 50 million years ago, for example, the Earth was around 10°C warmer than today.

Ice cores give us a detailed temperature record back for about 800,000 years (see figure 3). Over this period the climate has repeatedly shifted between ice ages - up to 8°C colder than today - and mild interludes similar to the last few thousand years.8 Warming and cooling have often happened rapidly, but there is no evidence that Earth has warmed as quickly as it is doing today for at least 50 million years.

FIGURE 3: ANCIENT AND MODERN



How the Earth is warming

Air bubbles in Antarctic ice reveal that temperature has tracked CO₂ closely for hundreds of thousands of years. Modern measurements show that CO₂ is now far above this prehistoric range.

(Source: Adapted from Woods Hole Research Center.)

Climate change varies strongly from place to place. Some of the greatest warming is far inland and near the poles.

(Source: IPCC Fifth Assessment Report)

More evidence for climate change:

- Ice, snow and permafrost is melting
- AS WELL AS WARMING Sea level is rising
 - There are more heatwaves and heavy rainstorms
 - The growing season in northern latitudes is becoming longer
 - The geographical ranges of species and diseases are changing

An increase of 0.85°C seems small compared with seasonal and daily variations, but it is already enough to make extreme high temperatures much more frequent. Mean summer temperatures that would have been anomalously high in the middle of the last century (occurring on average less than 0.2% of the time in a given location) became quite commonplace by the period 2006 to 2011, happening 4 to 13% of the time.⁹ As well as shifting the range of variation to higher temperatures, climate change may also increase the severity of these variations.10

What we have learned recently

- 1. The Arctic may be warming faster than ever. Over the past century, the Arctic has warmed at about twice the global average rate. For the decade from 2000 to 2010, one study shows a rate three to four times the global average.¹¹
- 2. The Southern Hemisphere is also warming. Previous temperature data had been much sparser than for the north, and showed no evidence of warming over Antarctica. New data shows that that the Southern Hemisphere is definitely warming, including most of continental Antarctica.12

at an alarming rate.

Ice, snow and permafrost are melting

Overview

Before the Industrial Revolution, climate varied in response to natural events such as volcanic activity and changes in the amount of energy arriving from the sun. These things still play their part today. Several volcanic eruptions during the 20th century appear to have had a brief cooling effect, for example. Natural factors that affect the climate come in two main categories – external factors and internal variability.

External factors change the amount of heat entering or leaving the climate system:

- The activity of the sun affects the amount of solar radiation that reaches Earth (see section 3a).
- Cycles in our orbit change the amount of sunlight reaching Earth over millennia.
- · Volcanic eruptions add particles to the atmosphere that block sunlight and can cool the Earth substantially for a year or two.
- Natural changes to the Earth's carbon cycle from volcanic eruptions and shifts in ocean circulation, for example - alter the concentrations of CO₂ and methane in the atmosphere.

Internal variability moves heat around within the climate system:

• Climate patterns such as the El Niño/La Niña-Southern Oscillation can move heat between ocean and atmosphere, affecting air temperature over periods of up to a few decades (see 'What caused the hiatus', page 11).13

All of these factors operate on different timescales and many have different regional impacts. However, climate models consistently find that natural factors alone cannot explain the warming seen over the past half century¹⁴ (see figure 4). According to the IPCC's 2013 report, it is "likely" (more than 66% probability) that natural external factors have affected global temperature by between -0.1 and +0.1°C since 1951, and the same for internal variability.¹⁵ The remaining warming over that period is "extremely likely" to have come from human activity (see section 4; and for the IPCC's definitions of likelihood see Appendix, page 42).

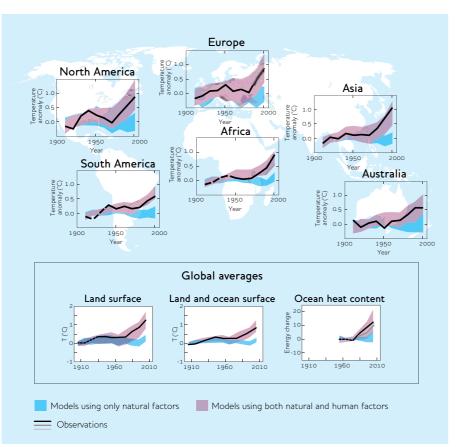
---- FAST FACT -----The climate changed in the past but never before this quickly



3. THE ROLE OF NATURAL FACTORS



FIGURE 4: FINGERPRINTS





Volcanoes disgorge fine particles into the the atmosphere that can cool the planet for a year or two.

Blue shading marks the range of climate model results for simulated

worlds where only natural factors

affect temperature. Purple marks

models in which both natural and

human factors are included. The black lines mark observations. This shows

that the pattern of warming in time

and place does not match natural

factors alone – human emissions

(Sources: IPCC Fourth and Fifth

must be included.

Assessment Reports)

WHAT CAUSED THE HIATUS?

After a period of rapid temperature rise from around 1970, warming of the atmosphere slowed down or stopped over the last 15 years. From 1998 to 2012, the warming trend was between -0.05 and +0.15°C per decade.¹⁶

Because of internal variability in the climate system, such decade-long pauses are to be expected.¹⁷ The early and mid-20th century even saw periods of cooling.

In their AR5 report, the IPCC attributed the recent hiatus to a **combination of three things** – internal variability, cooling by airborne particles from volcanic eruptions, and the downward phase of the solar cycle.¹⁸ However, they gave *low confidence* to the contribution from particles, and the solar effect is small (see section 3a).

The latest research suggests just one main cause: heat being absorbed by the ocean. Several models now come to this conclusion.¹⁹ For example, a model that incorporates observed water temperatures in the Pacific can reproduce the hiatus in global air temperature.²⁰ It also fits the regional pattern of measured temperature changes, and the seasonal pattern (while global winter temperatures fell from 1998 to 2012, summer temperatures continued to rise).

So the hiatus was probably caused by cooler water coming to the surface in the tropical Pacific and absorbing heat from the atmosphere. This is very similar to the well-known La Niña cooling pattern, but longer lasting. Climate models do generate such sea-surface fluctuations, but as yet they are unable to forecast *when* such phenomena will happen – a block to decade-scale climate forecasting.

Models suggest that during such pauses, most of the heat sucked from the atmosphere is carried down to the deep ocean below 300 metres.²¹ Ocean temperature measurements are not yet precise enough to confirm this for the period 1998 to 2012, although they do show that the heat content of the oceans continued to increase – so the planet as a whole was still warming.

If this picture is right, then when the tropical Pacific switches back to a warm surface state, global warming will accelerate.

What we still need to learn

1. Clouds are a large source of uncertainty. Clouds reflect sunlight, trap heat from below, and radiate their own heat to space. The net effect of low-level clouds is to cool the planet, while high-level clouds have a warming effect. The question is how clouds will be affected as temperature rises: will they amplify or dampen global warming? Models have not been very effective at capturing individual cloud formation, mainly because it happens on scales too fine for these simulations to handle explicitly. Particles from human and natural sources also affect the formation and properties of clouds by uncertain amounts.²² But the match between models and observations is improving,²³ and some recent research indicates that higher temperatures may reduce the amount of low-level cloud-cover - which would amplify warming.24

---- FAST FACT -----Changes in the sun cannot explain recent warming

2. Internal variability is not fully understood. Some natural fluctuations in weather and ocean circulation could have an appreciable effect over periods of a few decades,²⁵ and one study suggests that internal variability may have been responsible for as much as a third of the late-20th century warming.²⁶ Global warming may also change the pattern of natural climate cycles.²⁷

The role of solar variation

The sun is not constant. Over the past four billion years it has grown about 30% brighter, and its light output varies on shorter timescales as well. However, there is no evidence that variations in the sun could be responsible for the warming of the past few decades.

Since 1978, researchers have used satellites to reliably measure the intensity of sunlight reaching Earth. They see that over an 11-year cycle the sun regularly brightens and fades again.²⁸ The overall intensity of sunlight changes by about 0.1% during each cycle, enough to have some effect on global temperature. ²⁹ A fading sun has probably made a small contribution to the recent warming hiatus. But with a regular 11-year cycle up and down, this variation cannot account for the global warming observed over several decades.

Some researchers have claimed that there may be a longer-term upward trend in solar output over recent centuries. A study in 1995 suggested that such a trend could account for half of the observed global warming since 1860.³⁰ However, this calculation was based on assumptions that have now largely been refuted (for example, observations of other sun-like stars seemed to show that they could be in high and low states of activity, but more recent data show that these are in fact two chemically different types of star).

A more recent physical model of the sun's magnetic activity suggests a very gradual brightening by about 0.1% since the 18th century,31 which would be enough to raise global temperature by no more than about 0.15°C.32

And over the past few decades, satellite observations show no sign of an increase in solar brightness between cycles.33

Less obvious effects of solar variation could influence our climate:

- During each solar cycle, the sun's output of ultraviolet radiation can change by about 1.3%. In the fading phase of the last cycle, from 2004 to 2007, the decline in UV was even greater.³⁴ Ultraviolet heats the stratosphere and affects the pattern of high-altitude winds. The effect on global temperature is not clear, however.
- Energetic particles from deep space called cosmic rays may help to seed clouds on Earth.³⁵ When the sun is at a high point in its cycle, its magnetic field shields Earth from cosmic rays. In theory that could reduce cloud cover and warm the planet. But since the 1950s there has actually been a slight upward trend in cosmic rays detected on Earth, so this mechanism does not seem to be contributing to global warming.

4. THE ROLE OF HUMAN ACTIVITY

Overview

Climate scientists agree that the main factor behind global warming is human emission of greenhouse gases (GHGs) - primarily carbon dioxide, CO₂. The level of CO₂ in the atmosphere rose from about 280 parts per million (ppm) during pre-industrial times to 393 ppm in 2011. For a week in May 2013 it exceeded 400 ppm. It is now higher than at any time in at least the past 800,000 years. The increase is also much faster than in previous episodes of climate change.³⁶ Bubbles of ancient air trapped in ice cores show that at the end of each ice age,

CO₂ rose by only about 80 ppm over a period of 5,000 years³⁷, which in prehistoric terms amounts to a sudden climatic transition.

All the available evidence strongly suggests that this rise in greenhouse gases has caused the majority of the recent rise in global temperature. First, climate models can match the warming observed over the past century only by including the effects of greenhouse gas emissions (see figure 4). In addition, the geographical pattern of warming matches that predicted from an increase in greenhouse gases, as does the fact that the stratosphere is cooling, and that nighttime temperature is rising faster than daytime. Recent research has reinforced the conclusion that human emissions of GHGs are responsible for most of the warming observed in recent decades.³⁸ According to the IPCC's 2013 report, "It is extremely likely (more than 95% probability) that human influence has been the dominant cause of the observed warming since the mid-20th century".39



EFFECT

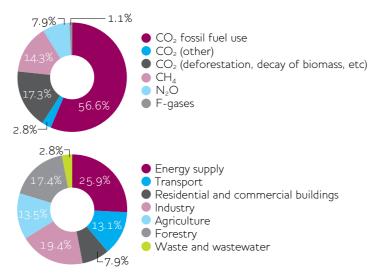
INDIRECT



Human greenhouse gas emissions are responsible for most of the global warming of recent decades.

---- FAST FACT -----Carbon dioxide levels have increased by about in the past 2 centuries

Greenhouse gases (GHGs) are generated by various natural and human sources (see figure below). They trap infrared radiation emitted by Earth's surface and atmosphere. That reduces the flow of heat to space and therefore warms the planet.



Source: Adapted from IPCC Fourth Assessment Report

GREENHOUSE GASES

Long-lived GHGs persist in the atmosphere for years to millennia, continuing to warm the climate long after they are first emitted:

- Carbon dioxide levels have increased by about 40% in the past two centuries, mainly due to the use of fossil fuels and deforestation. Some of the CO₂ being emitted now will remain in the atmosphere for many centuries.
- Methane is generated by agriculture (especially livestock), fossil fuel production, and the disposal of waste in landfill, as well as by natural sources including wetlands and oceans. Methane is a much more powerful greenhouse gas than CO₂, but is lost from the atmosphere in a few decades.
- Nitrous oxide is emitted by artificial fertilisers, and also in small quantities from fossil fuel combustion.
- Halocarbons (eg, chlorofluorocarbons) are the most powerful GHGs and some persist for millennia, but they are present in very small quantities. They were used as refrigerants and in other products and industrial processes before they were found to damage the ozone layer in the upper atmosphere, and were regulated internationally.

Short-lived GHGs persist from days to months:

- Tropospheric ozone (ozone in the lower atmosphere) is created by chemical reactions with carbon monoxide, nitrogen oxides, methane and other gases, all of which are emitted during human activities.
- Water vapour is the most abundant GHG in the Earth's atmosphere. Human activities have only a small direct influence on levels of water vapour, but a large indirect influence because rising temperature increases the atmosphere's capacity to hold moisture. That means water vapour acts as an internal feedback to amplify climate change. It also exerts more complex effects by forming clouds (see section 3).

As well as our GHG emissions, human activity can affect the climate through:

- Black carbon, or soot, from burning vegetation, wood and fossil fuels. It absorbs sunlight and so heats the atmosphere.
- Sulphate particles, from industrial pollution. These reflect sunlight and cool the atmosphere.
- Changing reflectiveness of Earth's surface. For example, replacing forests with agricultural land or desert increases the amount of sunlight reflected.

What we have learned recently

GASES

GREENHOUSE

OTHER HUMAN EFFECTS

Black carbon plays a greater role in warming than we thought. A new study estimates that soot emitted by fires and fossil fuels is the second most important human influence on the climate, after CO₂, and says that many models underestimate the effect of black carbon by a factor of three.⁴⁰

5. WHAT THE FUTURE HOLDS

5a. GREENHOUSE GAS EMISSIONS

Overview

Four different futures are used for climate projections in the AR5. These "representative concentration pathways" (RCPs) cover a range of possibilities for population, technology and economic development.⁴¹

At the high end, RCP 8.5 reflects "business as usual": rapid population growth and slow adoption of new technologies, resulting in high carbon emissions. By contrast, RCP 2.6 assumes very strong mitigation, with emissions falling rapidly after a peak in the early 21st century (see figure 6).42 Practically speaking, we can call these worst-case and best-case outcomes, respectively.

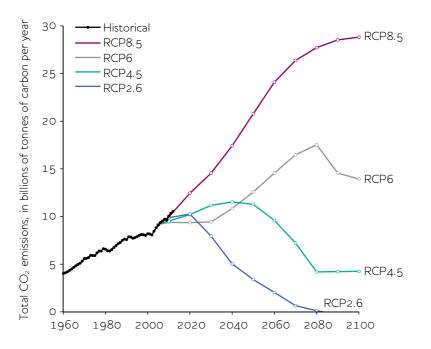


FIGURE 5: PATHS FOR THE FUTURE

Peak temperature will depend more on cumulative emissions than their timing. To have a 66% chance of staying within a 2°C warming target, the cumulative global emissions budget is estimated at 1,000 billion tonnes of carbon.43 (That is considering only CO, emissions. If methane and other more powerful greenhouse gases are included, the budget goes down to 800 billion tonnes or less). About 530 billion tonnes had already been emitted by 2011.



NOW	24.00	
NOW	2100	
In 2011, the mean global	In 2100, gas concent	
concentrations of the main	the various pathways	
greenhouses gases were	range from:45	
measured at:44	• CO ₂ : 421 ppm (for	
• CO ₂ : 391 parts per million (ppm)	936 ppm (for RCP	
• Methane: 1,803 parts per billion	• Methane: 1,254 t	
(ppb)	• Nitrous oxide: 344	

Nitrous oxide: 324 ppb



used in the latest climate models, ranging from the high end "business as usual" RCP 8.5, right down to a future of aggressive mitigation, RCP 2.6, in which net CO₂ emissions actually go negative by the end of the century. (The International Energy Authority [IEA] uses a different set of emissions scenarios running to 2035. The IEA 450 scenario is close to RCP 2.6. IEA New Policies is fairly similar to RCP 4.5, while IEA Current Policies is about half way between RCP 6 and RCP 8.5.)

Four future emissions pathways are

(Source: Adapted from Glen Peters/ CICERO)

trations set by ys used in AR5

or RCP 2.6) to P 8.5) to 3,751 ppb • Nitrous oxide: 344 to 435 ppb

> Approximately 530 billion tonnes of carbon had already been emitted by 2011.

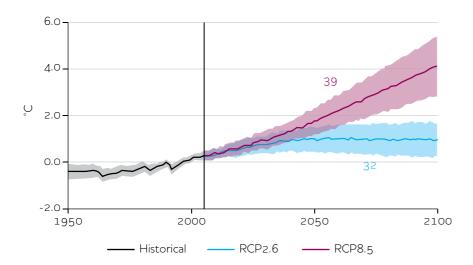
5b. WARMING

Overview

Because the oceans act as a buffer, the atmosphere does not respond instantly to increasing greenhouse gases; instead there is a time-lag in atmospheric warming. This means that greenhouse gases already emitted by human activity have locked the planet into 0.3 to 0.7°C of warming, on top of the 0.85°C increase we have already experienced.46

Climate models predict that if emissions go on unabated (the RCP 8.5 scenario), then global temperature is likely to rise by 2.6 to 4.8°C during the 21st century, reaching 3.2 to 5.4°C above pre-industrial levels. For the strong mitigation RCP 2.6 scenario, models predict a likely rise of 0.3 to 1.7°C in the 21st century, reaching 0.9 to 2.3°C above pre-industrial.47

FIGURE 6: PROJECTED TEMPERATURE CHANGE



Projections of global temperature change under high and low emissions. Shading shows the range of model outputs for each pathway, and the number of models involved is given next to each graph.

(Source: IPCC Fifth Assessment Report.)

Many experts argue that warming greater than 2°C above pre-industrial levels would pose a high risk of causing dangerous interference to the Earth's climate.48

MODELS OF EARTH

To predict what might happen to Earth's climate in the future, scientists build computer models that simulate the atmosphere, oceans and land surface.

These models are based on the known physics of:

- How air and water behave under the effects of gravity, pressure and heat flow, while sandwiched to the surface of a rotating planet
- How each gas in the atmosphere absorbs and emits different wavelengths of light

Global Climate Models (GCMs) cover the whole atmosphere and ocean. No existing computer could track every eddy and gust of wind, so to simplify things the atmosphere and ocean are pixelated: divided up into threedimensional grids. Atmosphere grid boxes are about a hundred kilometres across and a few hundred metres deep in state-of-the-art global models. Regional models have higher resolution.

Each property of the air or water inside a box, such as pressure and temperature, is represented with a single number. The model then calculates how neighbouring boxes affect one another, to predict how air and water move and how their temperature and other properties change. In the short term, that is what we call weather. Climate models do not need to predict what the weather will be on 7 February 2033; instead their aim is to predict the average weather over very long periods.

To add in the carbon cycle, a GCM can be coupled with models of vegetation and soils. That makes an Earth System Model, such as HADGEM-2, developed at the UK Meteorological Office's Hadley research centre.⁴⁹ There are several such models, and each one is run many times over to gauge the range of natural variability in the climate.



Some processes, such as convection and the growth of clouds, happen on scales smaller than a GCM's grid. Each of these is handled with a separate physical model called a parameterisation. For example, the rate of evaporation from the sea surface is calculated from local humidity, temperature and wind speed, according to an approximate formula. Such formulas can be checked and improved by comparison with field observations, but parameterisations remain a source of uncertainty within climate models.

The spread of possibilities for future climate can be judged by comparing the results of several different models, although only to a certain extent, because they are not entirely independent. Many make similar physical assumptions, and some share sections of computer code.⁵⁰

Models may all be neglecting some factors that affect climate, but their results have been tested. This is done by running a model for the conditions of the 20th century and comparing its output with actual observations. According to the IPCC "there is very high confidence that climate models reproduce the observed large-scale patterns and multi-decadal trends in surface temperature".51

---- FAST FACT -----By the end of the 21st century global temperature could have risen by up to above pre-industrial levels

What we still need to learn

- 1. The climate's response to rising CO₂ depends on some processes that are not fully understood, especially feedback from clouds and atmospheric particles (section 3, 'The role of natural factors'). Satellite observations are beginning to constrain these feedbacks, but a longer record of data is needed.52
- 2. Even if all processes were fully understood the picture is still blurred by internal variability of the climate: short-term jitter that makes it difficult to judge precisely how sensitive⁵³ the climate is to CO₂. This source of uncertainty will be reduced only gradually over time, as researchers build more detailed models of the oceans and other systems, and are able to compare longer model runs with observations.

5c. SNOW, ICE AND FROZEN GROUND

Overview

The Earth's snow, ice and frozen ground, together known as the cryosphere, make up the second largest component of the climate system after the oceans in terms of mass and capacity to store heat.54

The cryosphere is one of the most sensitive indicators of climate change. Worldwide, glaciers are retreating, and the great ice sheets of Greenland and Antarctica are slowly shrinking. Greenland alone is losing more than 270 billion tonnes of ice per year.55 The most dramatic change is in the Arctic (see figure 7), where over the last 30 to 40 years 60 to 75% of sea ice volume has disappeared,⁵⁶ and data indicates that the ice has not retreated this far for at least 1,400 years.⁵⁷ The IPCC are now confident enough to pin this at least partly on human activity, which they say is "very likely to have contributed to Arctic sea ice loss since 1979".58

FIGURE 7: SHRINKING ARCTIC SEA-ICE





Later this century, sea ice will probably vanish from the Arctic summer. Glacier melting will accelerate, raising sea level. Permafrost will thaw in some areas, releasing more greenhouse gases.

CLIMATE CHANGE AFFECTS THE CRYOSPHERE:

- **RISING TEMPERATURE**
- This can melt snow and ice and thaw frozen ground. INCREASED RAIN OR SNOW
- Heavier precipitation from a warmer climate and changed atmospheric circulation can lead to local increases in snow or ice.
- CHANGING AIR AND OCEAN CIRCULATION This can bring warmer water into contact with sea ice and with ice shelves, which normally help to restrain the flow of glaciers into the sea.59
- **BLACK CARBON (SOOT)**

CAUSES

Soot from coal and diesel combustion and agricultural fires⁶⁰ can be deposited on snow where it absorbs solar radiation as heat, and so accelerates snowmelt.⁶¹ This effect is greater in the Arctic than the Antarctic, which is more remote from sources of black carbon.62

Sea ice coverage in the Arctic, measured at its summer minimum in September each year by satellite observation. Although some years show an apparent 'recovery' the longer-term trend is clearly downwards.

---- FAST FACT -----

In the past few

decades,

60-75%

of Arctic sea ice has

disappeared

(Source: Adapted from National Snow and Ice Data Center)

Greenland's ice sheets are shrinking at a rate of more than 270 billion tonnes of ice per year.

THE THAWING CRYOSPHERE AFFECTS HUMAN ACTIVITY:

SEA-LEVEL RISE

Meltwater from the Greenland and Antarctic ice sheets is adding about 0.6 millimetres per year to sea level.63 This could accelerate in the future and cause a very large rise in sea level (see section 6a, 'Sea-level rise').

GROUND DESTABILISATION

Faster and earlier thaws can damage infrastructure such as ice roads, oil and gas pipelines and drilling rigs that rely on seasonally frozen land, rivers and lakes.

ECOLOGICAL CHANGE

Organisms dependent on sea ice, such as the polar bear, are among the most directly affected, and tundra ecosystems may also be threatened.

FRESH WATER AVAILABILITY

Runoff from ice and seasonal snowmelt is a vital source of water for human use and for fresh-water ecosystems.⁶⁴ Rising temperatures can cause this runoff to occur earlier in the season and more quickly, reducing the flow of fresh water later in the summer. Melting glaciers will initially provide more water, sometimes causing floods; then as they diminish so will the amount of water they provide.

EASIER ACCESS TO THE ARCTIC

The melting of sea ice is opening new shipping routes in the Arctic as well as easing access to natural resources such as oil and gas.

HYDROELECTRIC ENERGY New opportunities for hydroelectric energy will appear in the USA, Canada, Russia, Scandinavia and Greenland.65

THE THAWING CRYOSPHERE ALSO CREATES FEEDBACKS THAT CAN AMPLIFY CLIMATE CHANGE:

REFLECTION

Snow and ice reflect sunlight back into space, helping to keep the climate cool. As snow and ice melt, darker land or sea is exposed and absorbs more heat, amplifying global warming. This feedback is thought to play a central role in the high warming rate in the Arctic, where temperatures have risen at least twice as fast as the global average in recent decades.66

GASES RELEASED FROM PERMAFROST Evidence from past climates and recent observations suggest that thawing permafrost can release CO₂ and methane into the atmosphere, potentially causing rapid and extreme warming (see Section 6h: tipping points).67

NOW

- The Greenland and Antarctic ice sheets are shrinking, as are ice caps and glaciers around the world.
- Arctic sea ice is retreating.
- Antarctic sea ice is advancing in some areas, retreating in others.
- Permafrost is thawing in some areas.
- Seasonal snow cover is shrinking.
- though much less quickly. • Melting of the Greenland and West Antarctic ice sheets may become all-but irreversible.
- Permafrost area is virtually models project a global loss of roughly 80%.69

What we have learned recently

- 1. Thawing permafrost and melting ice is already releasing methane into the atmosphere (see Section 5j, 'Tipping points').⁷⁰
- 2. Advancing sea ice in parts of Antarctica is probably caused by stronger winds blowing from the land, pushing existing ice out to sea and exposing more water to freeze.71
- 3. By contrast, Antarctic ice shelves are weakening. These thick, floating shelves of ice act to restrain glaciers flowing from the interior of Antarctica, but relatively warm water is thinning them from beneath, leading to accelerating glaciers.72

What we still need to learn

The complex flow of ice within glaciers will determine the fate of the Greenland and West Antarctic ice sheets, which could potentially add several metres to global sea level (see section 6a, 'Hazards: sea-level rise' and section 6g, 'Tipping points'.) Ice dynamics is not well understood, although new research is improving the situation. For example, satellite observations of Antarctica show a network of glaciers stretching thousands of kilometres inland, demonstrating that ice is carried from the interior of the continent mainly through these channels - like tributaries of a river system - rather than by large-scale deformations of the ice sheet.⁷³ Sub-ice networks of water are thought to lubricate the flow of these glaciers.

FEEDBACKS

EFFECTS

2100

• Summer sea ice in the Arctic is projected to disappear entirely or almost entirely unless greenhouse gas emissions are strongly mitigated. Only in RCP 2.6 (the "best case" emissions pathway) is a substantial amount of sea ice projected to remain.68 Antarctic sea ice will retreat too, certain to shrink. Under RCP 8.5,

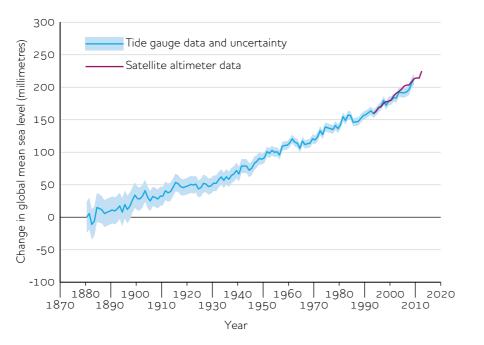
6. HAZARDS

6a. SEA-LEVEL RISE

Overview

Global sea level is currently rising at more than three millimetres per year (see figure 8). The scale and speed of future rises is debated, but some recent estimates are that sea level could rise by as much as 1.5 metres during the 21st century if the increase in greenhouse gas emissions is not abated.74

FIGURE 8: HIGH SEAS



Satellites confirm what tide gauges were already showing: sea level is rising, and rising faster in the most recent decades.

(Source: Adapted from Commonwealth Scientific and Industrial Research Organisation.)

> Sea level rises mainly because of melting land ice (which flows as water into the oceans) and thermal expansion (ocean water expands as it warms).

Predictions of sea-evel rise use one of two main methods:

- **Process-based models** calculate the thermal expansion of the sea and the melting of glaciers to add up all the separate contributions.
- Semi-empirical models seek a relationship between sea level and temperature from past observations. This approach produces higher sea-level projections than process-based models.

The IPCC's new report bases its estimates mainly on process-based models, and projects a global average sea level rise of 0.53 to 0.97 metres for unmitigated emissions (RCP 8.5), and 0.28 to 0.60 metres for strong mitigation (RCP 2.6).75 (All these figures are compared with average sea level between 1986 and 2005). Process-based models have improved since the 2007 AR4 report, but they still underestimate historical observed sea-level rise over the 20th century, so they may be giving conservative estimates of future rise.

According to semi-empirical models, sea level by 2100 could be appreciably higher. For RCP 8.5, estimates range from about 0.7 to 1.65 metres.⁷⁶ Beyond 2100 sea level will rise for centuries even if emissions are stabilised,

because of greenhouse gases already in the atmosphere and heat stored in the ocean, which will continue to melt the ice sheets.77

Currently, melting ice sheets are making a moderate contribution to sea-level rise, but they could reach a point where collapse is relatively rapid - perhaps on a timescale of centuries or less. Or they may melt slowly but irreversibly, locking in a long-term rise in sea level⁷⁸ (see 'Tipping points', page 33). This cannot be predicted with confidence because the internal dynamics of glaciers and ice sheets are not well understood.79

Greenland holds enough ice to raise global sea level by seven metres; the West Antarctic ice sheet could add three to five metres.⁸⁰

The West Antarctic ice sheet may be especially vulnerable because much of it is grounded below sea level, meaning that water could flow in and undermine it. That could even become a concern this century. In the AR5, the IPCC mention that the collapse of marine-based sectors of the West Antarctic ice sheet could cause global sea level to rise substantially above their estimates: "There is medium confidence that this additional contribution would not exceed several tenths of a meter of sea level rise during the 21st century."81

A rise in sea levels could pose one of the greatest threats to populations, economies and infrastructure in low-lying island and coastal regions, especially as tropical storms may become more violent, generating larger storm surges (see section 6c, 'Hazards: storms and floods'). Cities and settlements in the river deltas of Africa and Asia are among the most vulnerable and least able to spend on defence and adaptation.⁸² Floods could also threaten major cities in the developed world such as New York and Miami. One 2013 study calculates that if flood defences are not improved, total economic losses could reach a trillion dollars per year by 2050.83

Sea level will rise more in some areas, less in others, because of changes in ocean currents and the shifting burden of ice. The ice sheets are so heavy that they depress the Earth's crust, and their gravity pulls water in around them – so as the ice melts both those effects will lessen. Some of the fastest-rising sea levels are expected along the US east coast and around the equator.84

What we have learned recently

Groundwater pumped out for human supplies also adds to sea-level rise. Previously considered negligible, it may have contributed 13 to 25% of observed sea-level rise according to two studies.⁸⁵ Groundwater mining could add between five and 20 centimetres to sea levels in the 21st century.86

---- FAST FACT -----By the end of the century sea level could have risen by up to 1.5

metres

6b. HEATWAVES AND DROUGHT

Overview

As the world warms, it is virtually certain that temperatures considered extreme today will become more frequent, while extreme cold will become rarer.

Trends in extreme weather are hard to measure, as these are rare events and most regions lack long-term records, but the number of extremely hot days has increased since 1950.87 Research suggests that over the past few decades, heatwaves – very high temperatures lasting several days – have become more frequent.⁸⁸ The IPCC considers heatwaves very likely to become more frequent and longer lasting this century.89

Some regions have also seen longer and more intense droughts, especially southern Europe and West Africa.⁹⁰ Drought depends on rainfall as well as temperature, making it more complex to predict, but total precipitation is expected to change in a way that enhances regional extremes. So already dry areas such as the subtropics will probably see more droughts.91

Even in the "best case" scenario RCP 2.6, both heatwave and drought risks increase. The higher the emissions pathway, the greater the risk.

Because there is a limit to how far humans can physiologically adapt to heat stress, unmitigated climate change could eventually result in some regions becoming uninhabitable,⁹² although only for global warming of more than 7°C.

HEAT HAZARDS	 Drought can be caused when higher temperatures lead to more evaporation, or if shorter winters or earlier snowmelt reduce the summertime water supply. In turn, drought can prevent plants from cooling the air, increasing the intensity of a heatwave. Hot weather may become prolonged more often, because as polar regions warm up, the jetstream is expected to meander more slowly, blocking the movement of weather systems. (The same goes for cold and wet weather.)
EFFECTS	 Heat-related deaths from cardiovascular and respiratory illnesses will rise, while deaths from cold will fall. Drought's main threat is to agriculture. Even modest warming of 2°C above preindustrial levels would expose about 8% of the world's population to new or aggravated water scarcity, according to a recent study (and see 'Hazards: food and water').⁹³ Extreme heatwaves can damage crops even in the absence of drought.⁹⁴ Drought can also affect power generation because river water is used for cooling nuclear and fossil-fuel powerplants, as well as driving hydroelectric plants (see section 6f, Hazards: food and water).

What we have learned recently

Researchers have improved their ability to attribute individual events to warming. Although it is not possible to say with certainty that a particular event such as the 2003 heat wave in Europe was caused by climate change, studies have calculated the probability that certain events will occur for a given climate, and concluded that some extremes would have been highly improbable without greenhousegas induced warming.95

6c. STORMS AND FLOODS

Overview

Warm air can hold more moisture, so in a warming world extreme rainstorms are almost certain to become more intense. Downpours have already become more frequent in some regions.96

Precipitation becomes a serious problem when it overwhelms river systems and leads to flooding. This also depends on other factors, including soil moisture and changing land use, making prediction complex. But models project that over the coming decades already wet regions, especially near the equator, will see more downpours and more floods.97

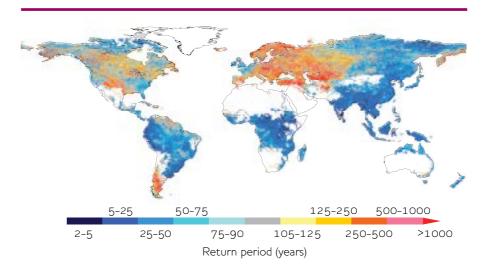
According to one study, increases in extreme precipitation could be underestimated by most models when compared with observations.98 The intensity of hurricanes and other tropical cyclones is likely to be boosted by increased evaporation from warmer seas. Powerful category four and five storms are already becoming more common, while weaker storms become less common, according to recent research. One study concludes that "the proportion of Category 4 and 5 hurricanes has increased at a rate of 25 to 30% per °C of global warming". In the future, the IPCC concludes that maximum wind speeds and precipitation rates are likely to increase further.99

Coastal flooding is expected to increase, mainly because of a combination of rising mean sea levels and higher storm surges as cyclones become more intense.100

Material damage from extreme weather increased eight-fold between the 1960s and the 1990s in inflation-adjusted terms, a faster increase than population or economic growth.¹⁰¹

What we have learned recently

Flooding is predicted to increase most in southeast Asia, peninsular India, East Africa and the northern Andes, according to a new study that compares 11 different climate models.¹⁰² The same study predicts a decrease in flooding across most of North America, Europe and central Asia. The overall effect is to expose a larger population to flooding: under the RCP 8.5 "worst case" scenario it goes up from around 20 million people today to about 100 million in 2100.



---- FAST FACT -----Powerful Category 4 and 5 storms are already becoming more common

Flooding is projected to become more frequent (blue) in a majority of areas, but rarer (warm colours) in others, if greenhouse gas emissions are unmitigated (RCP 8.5 - the "worstcase" pathway). For each location, this map refers to a level of flood that would have been a once-in-100-year event in the climate of the late 20th century. It shows how that expected return period is projected to change by the period 2071 to 2100. The darkest blue marks places where such floods are projected to happen on average every two to five years.

(Source: Hirabayashi et al., "Global flood risk under climate change," Nature Climate Change, 3, pp. 816-821 [2013])

6d. EXTREME WAVES

Overview

AND WAVE

MIND

The largest ocean waves can threaten ships, drilling platforms and wind turbines. In many regions these extreme waves are growing in size. Models project that in some regions global warming will probably increase wave height further, especially in some crucial shipping areas such as the North Atlantic.

- Waves are built up by wind blowing over the sea surface. Local wave heights will increase if wind speeds increase, or if the wind blows uninterrupted for a longer time or over a greater distance.
- **Rolling waves called swells** can travel thousands of kilometres and boost the size of locally generated waves. Swells generated by storms in the Southern ocean are expected to increase in size, and add to wave height as far away as the North Atlantic.¹⁰³
- **Significant wave height** captures wave conditions in a single number, roughly equal to the average trough-to-crest height of the highest third of waves seen at any one time and place. Ships are designed to withstand sea conditions rare enough to occur only once in 20 years, which in the North Atlantic means a significant wave height of at least 16 metres.¹⁰⁴
- Very rare "rogue waves" can reach more than twice the significant wave height, and tend to be very steep, making them even more dangerous. Warming may increase the number of rogue waves.105

Since the 19th century, wave height has been recorded systematically by visual observation from ships. More recently, fixed buoys, marine radar and satellites have measured wave height, and models have been used to calculate wave height retrospectively based on the weather conditions.

Each method has some shortcomings. Models are indirect. Ships steer clear of storms. Buoys sample only a few locations. The satellite record is brief, so it may not fully reflect long-term trends. They give differing numerical results, but all show that over the past half century extreme wave height has increased in the North Atlantic, the Southern Ocean and the Northern Pacific.¹⁰⁶

In the future, wave height is very likely to increase in the Arctic as global warming melts sea ice to expose more water. Elsewhere, future wave height is harder to predict as it depends on changing weather patterns, but a number of studies predict that by the end of the 21st century, extreme wave heights are likely to increase further in parts of the North Atlantic and northern Pacific. 107

For once-in-20-year sea conditions, some of the forecasts predict increases of up to two metres in significant wave height (see box). According to one study on the structural collapse of ships, a half-metre increase could raise the probability of failure by 50%.108

As well as posing a risk to shipping and offshore installations, larger waves could damage some shallow-water ecosystems,¹⁰⁹ affect coastal facilities such as ports, and increase coastal erosion.

6e. OCEAN ACIDIFICATION

Overview

Seawater has absorbed 25 to 30% of the CO₂ emitted by human activity.¹¹⁰ This has changed the ocean's chemistry. When CO₂ dissolves in water it forms an acid, and the acidity of the ocean has already increased by 30%.111

If carbon emissions continue unabated, models predict that by the end of the 21st century the ocean's acidity will be 200% higher than its preindustrial level.¹¹² Even in a mid-range emissions scenario that is predicted to be at least 100% higher, making the ocean more acidic that it has been for at least the past 20 million years.¹¹³ Ocean chemistry will probably not recover for tens of thousands of years.

In Earth's past, major episodes of ocean acidification have been accompanied by large-scale extinctions among plankton and other marine species, many of which adapt to a narrow range of acidity.

CO₂ emissions affect ocean acidity through two mechanisms: • Some CO₂ is absorbed by the ocean, where CO₂ reacts with water to form carbonic acid. Warming reduces the solubility of CO₂ in the ocean, which partially reduces the rate of ocean acidification.¹¹⁴ SEAS A more acidic ocean has consequences for: ACID

• Sea creatures that build their shells or skeletons from calcium carbonate, which will die above a certain level of acidity

- Sea grasses, which may benefit from higher levels of CO₂
- **Global warming,** which could be amplified if acidification reduces the emission of cooling sulphates from the ocean into the atmosphere.115



Coral reefs are built from calcium carbonate and so are under threat as the ocean becomes more acidic.

- FAST FACT -----The ocean's acidity has increased by

Many organisms including corals, plankton and molluscs build their shells and skeletons out of calcium carbonate.

As the ocean's acidity rises, such structures may begin to dissolve, with the effect differing in strength for different organisms.¹¹⁶ Many of these organisms are at the base of food chains, so acidification may threaten not just individual species but the ecosystems they support as well as related fisheries and tourism.

The Southern and Arctic Oceans and the Pacific coast of North America are especially vulnerable, as cold water can hold more CO₂. Along the Pacific coast of North America, for example, water acidic enough to dissolve calcium carbonate shells is already seen each spring time. Although this is a natural phenomenon, the range of acidic water is spreading.¹¹⁷

Even hitting tight emissions targets may not be enough to preserve many coral reefs. In October 2013, the International Programme on the State of the Ocean released a report concluding that "at CO₂ concentrations of 450–500 ppm (projected in 2030-2050), erosion will exceed calcification in the coral reef building process, resulting in the extinction of some species and decline in biodiversity overall."118

What we have learned recently

- 1. Parts of the ocean may become hostile to shell-building organisms earlier than anticipated. It had been thought that atmospheric concentrations above 600 ppm (a level that will be passed before the end of the century unless action is taken to reduce emissions) might render Southern Ocean and the Arctic Ocean unable to support some organisms that build carbonate shells. Recent studies find that this threshold could be crossed at a level of 450 ppm, which is likely to be reached by mid century.¹¹⁹
- 2. Ancient climate records are providing insight into past ocean acidification. In one review, researchers concluded that the current rate of CO₂ release could cause chemical changes in the ocean unparalleled in the past 300 million vears.120
- 3. Acidification could amplify global warming. Some marine organisms emit dimethyl sulphide, which goes on to form sulphate aerosol particles in the atmosphere, which reflect incoming sunlight. A new study concludes that in a more acidic ocean, organisms generate less dimethyl sulphide.¹²¹ By reducing our reflecting shield of sulphate particles, that would amplify climate change an effect not yet included in climate models.



Acidification also affects industries reliant on the ocean, such as fisheries and tourism.

6f. FOOD AND WATER

Overview

Climate change is already affecting water and food resources, and is expected to have an even greater impact over the coming century.¹²²

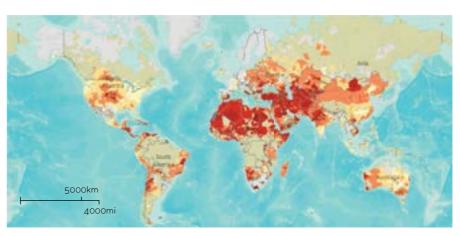
Drought is the greatest concern, with already dry regions likely to become drier.¹²³ In many models, much of the world becomes increasingly arid, including most of the Americas, Australia, southeast Asia, Africa, southern Europe and the Middle East. According to one study, people living in these regions may see a switch to persistent severe droughts in the next 20 to 50 years.124

Loss of mountain glaciers will reduce summertime river flow in some areas. For example, meltwater from the Himalaya forms an important summer supply to the Indus and Brahmaputra river basins,125 while Andean glaciers supply parts of South America.¹²⁶

Where total rainfall increases, it may move to the wrong season. That could be as damaging as drought, unless facilities can be built to capture and retain fresh water.127 An excess of rainfall can also wash away soils, waterlog crops or bury them under silt.

Higher temperatures will have a direct effect on crops, positive or negative depending on the location,128 with extreme heatwaves especially damaging.129 Warming of 1 to 3°C will probably increase crop yields slightly in some highlatitude regions, but significantly reduce them in lower latitudes. Warming above 3°C is projected to reduce the potential for global food production.¹³⁰

FIGURE 10: WATER STRESS, 2095



Exceptionally less stressed (<0.125x) Extremely less stressed (<0.125–0.357x) Significantly less stressed (<0.357-0.500x) Moderately less stressed (<0.500-0.588x) Wetter but still extremely high stress (<0.588 and baseline water stress (2095) >80%) Near normal conditions (<0.588–1.7x) Drier but still low stress (>1.7x and baseline water stress (2095) >10%)

32 | GLTE | Statkraft | DNV GL | Shell | TCS | Xyntéo

Moderately more stressed (1.7–2x) Severely more stressed (2–2.8x) Extremely more stressed (2.8–8x) Exceptionally more stressed (>8x) Missing data (no data) Uncertainty in magnitude Uncertainty in direction

Changes in water stress by 2095, in a high-emissions scenario similar to RCP 8.5. as assessed by the World Resources Institute. Under lower emissions scenarios, the changes are less severe.

(Source: Interactive maps available at www.wri.org/our-work/project/ aqueduct)

Climate change is expected to impair fresh water resources:131

- Warming will reduce the supply of fresh water in many regions.
- Shifting weather patterns, and the loss of glaciers and seasonal
- snowpack could change the timing of fresh water availability. Floods, rising sea level and increasing evaporation could contaminate fresh water resources with salt and other substances.

With consequences for:

- Crop yields. Increased atmospheric CO₂ can boost yields, but higher temperatures and drier soils tend to decrease them. Switching to less water-hungry crops can help, but the overall effect is expected to be negative for many parts of the world during the 21st century. More frequent floods, droughts, heat waves and forest fires are also liable to reduce agricultural production.¹³²
- Human health. For example, drought can lead to malnutrition and the spread of infectious diseases. The World Health Organisation estimates that the effects of climate change already claim 150,000 lives per year.133
- Aquatic ecosystems and fisheries, which are vulnerable to changes in temperature and chemistry of seas and rivers, as well as weakening ocean circulation.¹³⁴ In the Arctic, marine ecosystems will change as sea ice retreats,¹³⁵ because nutrientrich waters at the ice edge are a prime site for plankton growth; while ice loss may also promote increased photosynthesis. The net effect on fisheries is highly uncertain.
- Electricity generation. Many power plants use river water for cooling, and, by mid-century, lower river flow and higher water temperature could reduce generating capacity by 6.3 to 19% in Europe and 4.4 to 16% in the United States.¹³⁶ Hydroelectric plants depend on water flow and are vulnerable to climate change in some regions.¹³⁷ Hydraulic fracturing or 'fracking' for gas and oil production is also water intensive, as is carbon capture and storage.

And knock-on effects:

- Reduced food production, combined with increasing demand from population and economic growth, could drive up food prices 10 to 60% by 2030.138
- Worsening scarcity of water and food could cause human migration out of particularly stressed regions, and increase the likelihood of war.

What we have learned recently

Climate change is already lowering yields. A recent study calculates that the global production of maize and wheat between 1980 and 2008 was 3.8 and 5.5% lower, respectively, than it would have been in the absence of rising temperatures.139

6g. TIPPING POINTS

Overview

Abrupt climate change has happened in Earth's past,¹⁴⁰ and in theory it could happen again. Some element influencing climate, such as the amount of greenhouse gases in the atmosphere or the reflectivity of the Earth's surface, could cross a key threshold and trigger a transition to a new climate state.¹⁴¹ Such thresholds have been called 'tipping points'. The term usually implies that the transition would be relatively rapid - on a timescale of decades or less - and often that it would be effectively irreversible.142

Most climate scientists agree that if emissions continue unabated, some tipping points will be reached eventually, but there is little consensus and much uncertainty. It is especially unclear whether any of these events will occur soon.

In the AR5, the IPCC imply that it is unlikely we will hit a tipping point in the 21st century: "There is no evidence for global-scale tipping points in any of the most comprehensive models."143 They may be too conservative on this point, because so little is known about most tipping points. In particular, these models fail to capture the detailed processes underlying permafrost thaw and ice-sheet collapse.

Some recent research suggests we may be more likely to hit tipping points than was previously thought.144 One survey of climate scientists concluded that if global warming exceeds 4°C, hitting at least one tipping point will become more likely than not.145

Even if the probability is not so high, tipping points present a great risk. The consequences of collapsing ice sheets, drying Amazon or runaway permafrost thaw would be catastrophic.



---- FAST FACT -----

Warming has already lowered global production of maize wheat

Warming and deforestation could dry out the Amazon rainforest.

---- FAST FACT -----

The Amazon could rapidly shift to grassland if of existing forest is cut down

TIPPING POINTS

- Permafrost thaw. As Arctic soils thaw out, they can decompose and release methane and carbon dioxide, generating further warming. This is already happening in some parts of the Arctic. At some point, as yet unknown, this could become so serious that warming runs away, caught in a vicious circle of feedback. Permafrost varies in its properties from place to place and it has not been mapped in detail, making this process very difficult to model. However, many scientists believe that the permafrost contains very large amounts of carbon, making this perhaps the most serious of the potential tipping points. (And see under 'What we have learned recently').
- Ice sheet collapse. The West Antarctic ice sheet rests on ground below sea level, which may make it mechanically unstable, and susceptible to rapid collapse on a timescale as short as a few decades. In Greenland, melting could become irreversible because of two feedback mechanisms: exposed land or water absorb more heat than ice and snow, and glacier surfaces become warmer as melting reduces their elevation. The tipping threshold of the Greenland ice sheet could be very low (see below) – although it would probably take several centuries to melt.146
- Drying Amazon. Warming and deforestation could dry out the Amazon rainforest. Transpiration from the trees helps to maintain a damp local climate, so when too much of the forest is lost a tipping point could be reached. Some models predict up to 70% of the rainforest will be gone by the end of the century, replaced by savannah.¹⁴⁷ As well as the loss of biodiversity, this would release large amounts of CO₂, amplifying climate change.
- **Circulation shutdown.** A global conveyor belt of ocean currents is driven by the sinking of cold salty water in the North Atlantic. This process could slow or halt as temperatures rise and more fresh water is flushed into the North Atlantic from increased river flow and melting ice. If circulation is slowed or shut down, it will have many consequences for global climate including colder winters in Europe. Climate scientists generally agree that a shutdown is not likely to happen this century.
- Shifting monsoons. The Indian summer monsoon has changed its strength and variability in the past, and models suggest that it could suddenly weaken, leading to much more frequent droughts.¹⁴⁸ Predicting whether this will happen is difficult partly because the monsoon is subject to conflicting influences. Rising greenhouse gas concentrations and consequent warming may increase the intensity of the monsoon.¹⁴⁹ On the other hand, aerosol particles from industrial pollution reflect solar heat and are thought to weaken the monsoon – a process that may already be happening.¹⁵⁰ Meanwhile, some models predict that the West African monsoon could also be disrupted, which would increase precipitation over the Sahel and Sahara, leading to greening-a possible rare example of a positive climate tipping point.¹⁵¹

What we have learned recently

- 1. The Greenland ice sheet could be highly sensitive to warming. Earlier research put its threshold at 3.1°C above pre-industrial temperatures.¹⁵² According to a recent study, however, the ice sheet's eventual disappearance will become inevitable at a threshold somewhere between 0.8 and 3.2°C above pre-industrial levels.153 As 0.8°C of warming has already taken place, this threshold may already have been passed. In July 2012, NASA satellites observed melting over 97% of the ice sheet's surface.¹⁵⁴
- 2. Permafrost tipping points may be more serious than was thought.155 In the Fifth Assessment's highest warming scenario RCP 8.5, the Arctic will have warmed by about 7.5°C in 2100, which according to one study could cause permafrost to emit 380 billion tonnes of methane and carbon dioxide.156 A large area of especially carbon-rich permafrost in north-eastern Siberia contains an estimated 500 billion tonnes of carbon, and a recent study suggests that 9°C of local warming – possible this century given amplified warming in the Arctic - could release some three-quarters of its carbon over about a century.157 In 2010 researchers reported that another huge reservoir of carbon - permafrost buried under frozen sea beds off the north coast of Siberia – could be more vulnerable to thawing than land-based permafrost.¹⁵⁸ Methane is already seeping out from coal beds and natural gas deposits that had been trapped under glaciers and permafrost.¹⁵⁹
- 3. The Amazon may be more delicate than previously thought. New research concludes that a rapid shift from forest to grassland could happen if more than about 10% of the existing forest is cut down.¹⁶⁰
- 4. Some scientists think that Arctic sea ice passed a tipping point when it receded in 2007.161

7. THE FIFTH ASSESSMENT **REPORT OF THE IPCC –** THE PHYSICAL SCIENCE BASIS

Overview

In October 2013, the Intergovernmental Panel on Climate Change released their first full-scale survey for six years. 'Climate Change 2013: The Physical Science Basis', 162 part of the organisation's Fifth Assessment Report (AR5), is the result of work by several hundred scientists to contribute, review and combine existing climate research to produce a consensus.

In most respects, the conclusions of the AR5 are strikingly similar to those of the IPCC's last comprehensive report, the AR4, in 2007. While there are a few significant differences (see box), both reports conclude that the main cause of global warming over the past century is human emissions of greenhouse gases, and that warming and its hazardous side effects will accelerate over the next few decades. In general the conclusions of the AR5 are more confident than the AR4, being based on updated observations and more detailed models.

The range of projections for the future is roughly the same. That too is no great surprise, because the largest uncertainty about the coming century is not scientific but political - how much carbon dioxide and other greenhouse gases will we emit?163



Inputs:

- Observations of the whole climate system have improved since the AR4. They also encompass an extra six years of data, which is significant where more advanced technology such as satellite observation has only been recently deployed. For example, the ARGO network of ocean probes¹⁶⁴ was only completed in 2007.
- Models have higher resolution. They also give a more realistic treatment of the carbon cycle and the behaviour of vegetation, and include more components of the climate – more types of airborne particle, for example.¹⁶⁵ The understanding of several geophysical processes including glacial dynamics and the water cycle has also improved.

Conclusions:

AR4

AR5 VERSUS

- Both assessments attribute warming mainly to human emissions of greenhouse gases, but where the AR4 considered this conclusion "very likely" (having a probability of more than 90%), the AR5 is more confident, calling it "extremely likely" (more than 95%).
- Since the AR4, evidence of human influence has grown in warming of the atmosphere and the ocean, changes in the global water cycle, reductions in snow and ice, sea level rise, and changes in some climate extremes.
- The AR5 directly attributes much of the loss of Arctic sea ice to human activity.
- Projections of sea-level rise in the AR5 are higher, partly because the models now allow future increases in the flow of ice through the glaciers of Greenland and Antarctica.
- The AR5 is more confident that extreme high sea levels (storm surges on top of steadily rising seas) will occur more often by 2100 ("very likely" versus "likely" in the AR4).
- The AR5 sets out a cumulative global emissions budget of one trillion tonnes of carbon (see section 5a). Hitting that level would give us a roughly even chance of avoiding warming above 2°C.

By its nature, such a large group often has to err on the side of caution in order to reach a document that all participants can agree on. After the AR4, the IPCC was criticised for overstating the future melting of Himalayan glaciers - a mistake that was soon rectified, but gathered a lot of unwelcome media attention. So while the AR5 is by far the most authoritative and detailed description of the present state and possible future of the climate, some of its conclusions may be conservative.

This document includes a few differences with the AR5:

· We include a method of projecting sea-level rise that gives somewhat higher values than the one chosen by the IPCC (see section 6a, 'Hazards: sea-level rise'). With both, we can represent the full range of results coming out of climate science.

---- FAST FACT -----The main cause of global warming over the past century is human emissions of greenhouse gases

- Our assessment of tipping points is less optimistic than the AR5 consensus. The majority of climate scientists consider that Earth will probably not hit a major tipping point this century. But if it does, then the downside would be extremely serious - so we think it important to discuss the more pessimistic studies as well (section 6g).
- We also have access to research too recent to have been included in the AR5. For example, new studies pin the recent warming hiatus mainly on a cool phase of Pacific ocean circulation (see section 3, 'The role of natural factors'), where the AR5 gave a more open verdict.

8. FREQUENTLY ASKED QUESTIONS

Climate science is a dynamic field, involving debate and uncertainty (see, for example, 'What we still need to know' in the preceding sections). However, some well-established conclusions are often called into question by critics outside the climate-science community. Here we present some frequently asked questions to which the science already has firm or fairly firm answers.

Do we know the Earth is really warming?

High-quality global temperature measurements go back 150 years.¹⁶⁶ These records have been analysed carefully by independent research groups, and they show that average global surface temperatures have increased by about 0.9°C over that time (see section 2).

Has global warming stopped?

No, but the warming of the atmosphere has paused. The rise in global air temperature was slow or even zero between 1998 and 2012, probably because of a natural climate cycle bringing cooler water to the surface of the tropical Pacific (see section 3). Measurements show that the Earth as a whole continued to warm, with heat going into the oceans.

Is Arctic sea ice actually increasing?

There was more ice in summer 2013 than in 2012, but this is just short-term variation. The long-term trend is clearly and rapidly downwards (see section 5c). Even in summer 2013, the sea ice covered less area than in any satellite-mapped year before 2007.¹⁶⁷ The ice is also getting thinner.

Can warming be explained by changes in the output of the sun?

The sun does change in brightness, waxing and waning over an 11-year cycle enough to nudge global temperatures up and down by about 0.1°C, but there has been no long-term increase in brightness that could explain the warming trend of the past half century (see section 3).

How does warming today compare with past climate change?

It is extremely rapid. Earth's climate has always varied in response to natural causes, sometimes warmer and sometimes colder than today. Compared with many prehistoric changes, the climate change of the past century is modest in scale; but it is happening quickly (see section 2). In a historical context it is already



substantial: temperatures are probably higher now than at any time in the past two thousand years.¹⁶⁸

Did CO_2 drive warming in the past, or was it the other way around?

It was almost certainly both. Ice-core records show that around the end of each ice age, temperatures in Antarctica started to rise 600 to 1,000 years before the rise in CO_2 , implying that CO_2 did not trigger warming. That initial warming is thought to have been caused by cyclical changes in Earth's orbit, which match the sequence of ice ages. The slightly higher temperatures would then have triggered the release of CO_2 from the oceans, amplifying the initial change.¹⁶⁹ This release of CO_2 is thought to have been enough to cause about half of the warming that ended the ice ages.¹⁷⁰ (There is also evidence that the timing of temperature versus CO_2 is different outside Antarctica. A recent study using about 70 to 80 proxy records of temperatures from different parts of the planet suggests that CO_2 lagged temperature only very briefly, and after that it actually lead the rise in average global temperatures.¹⁷¹)

Are humans or natural events behind the recent increase in atmospheric CO₂?

Humans. Natural flows of CO₂ between land, sea and air are huge (hundreds of gigatonnes each year) but they are usually in balance. That is, the rate of flow into the atmosphere is about equal to the rate of absorption from the atmosphere.¹⁷² One useful analogy is to picture carbon flows like the water running into a bath (from a tap) and at the same time running out of the bath (down the drain). If the input and output are in balance, large amounts of water may be running through the system but the level in the bath stays the same. A relatively small addition of water may then tip the balance and cause the bath to overflow. In the same way, before the onset of the Industrial Revolution, ice-core records show that atmospheric concentrations of CO₂ stayed between 180 and 300 parts per million for at least the past 800,000 years. Now human CO₂ emissions (which reached a record 31.6 gigatonnes in 2011¹⁷³), mainly from fossil fuel combustion, have tipped the balance. As a consequence, atmospheric levels have increased with unprecedented speed, reaching 400 parts per million in some places during the spring of 2012. Once in the atmosphere, CO₂ can linger for many centuries, so CO₂ emissions from human activities accumulate over time. (See section 4.)

MORE EVIDENCE THAT THE RISE IN CO₂ IS BEING DRIVEN BY HUMAN ACTIVITY:

- The other two sources that might have put so much CO₂ into the atmosphere are the ocean and the Earth's vegetation, but measurements show that these are currently absorbing more than they emit.
- One heavier form of carbon, C-14, has decreased in the atmosphere. That indicates the new carbon has come from fossil fuels, which do not contain C-14, rather than biological sources, which do.
- The timing of the increase in atmospheric CO₂ coincides with the onset of the Industrial Revolution.

Is Antarctica warming too?

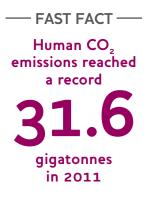
Yes. Based on limited observations, scientists previously believed that temperatures in the interior of the Antarctic continent were roughly steady, and the mass of the East Antarctic ice sheet might be increasing slightly. With more data, it has become clear that Antarctica is warming, and losing ice.¹⁷⁴ (See section 5c.)

Can the science be trusted?

Some errors were identified in the IPCC's 2007 Fourth Assessment Report, including an overestimate of the future melting rate of Himalayan glaciers. Those errors did not affect the report's main conclusions on the causes of warming during the last 150 years.

In 2009, emails were hacked from the University of East Anglia's Climate Research Unit (CRU). They included statements by climate scientists that were interpreted by some commentators to imply manipulation of the data. However, four independent investigations of the CRU cleared the climate scientists of any wrongdoing.¹⁷⁵

The large body of evidence showing the scale of warming, and its probable causes, has proved to be robust.



APPENDIX: IPCC AR5 LIKELIHOOD TERMINOLOGY

ENDNOTES

LIKELIHOOD OF 1	THE OCCURRENCE	OUTCOME
	THE OCCOMMENCE	

Virtually certain	> 99% probability	
Extremely likely	> 95% probability	
Very likely	> 90% probability	
Likely	> 66% probability	
More likely than not	> 50% probability	
About as likely as not	33 to 66% probability	
Unlikely	< 33% probability	
Very unlikely	< 10% probability	
Extremely unlikely	< 5% probability	
Exceptionally unlikely	< 1% probability	

- 1. Three main datasets of global surface temperatures go back to about 1850. They show similar results from month to month and all agree on the overall warming trend observed over the past century. These datasets are maintained by the UK's Met Office, the US Goddard Institute for Space Studies (a division of NASA) and the US National Climatic Data Center (a division of the National Oceanic and Atmospheric Administration). IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 2: "Observations: and Atmosphere and surface," available at www.climatechange2013.org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_ Chapter02.pdf.
- 2. Hansen et al., "Global surface temperature change," Review of Geophysics 48, RG4004 (2010); Screen et al., "Local and remote controls on observed Arctic warming," Geophysical Research Letters 39, L10709 (2012).
- 3. Easterling and Wehner, "Is the climate warming or cooling?" Geophysical Research Letters 36, L08706 (2009).
- IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical 4. Science Basis, Chapter 2: "Observations: Atmosphere and Surface," available at www.climatechange2013.org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_ Chapter02.pdf.
- 5. Levitus et al., "World ocean heat content and thermosteric sea level change (0-2000 m), 1955–2010," Geophysical Research Letters 39, L10603 (2012).
- Levitus et al. "World ocean heat content and thermosteric sea level change (0-2000 m), 6. 1955-2010," Geophysical Research Letters 39, L10603 (2012); Church et al., "Revisiting the Earth's sea-level and energy budgets from 1961 to 2008," Geophysical Research Letters 38, L18601 (2011).
- 7. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 3: "Observations: Ocean," available at www.climatechange2013. org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_Chapter03.pdf.
- Jouzel at al., "Orbital and Millennial Antarctic Climate Variability over the Past 800,000 8. years", Science 317, pp. 793-796 (2007).
- 9. Hansen et al., "Perception of climate change," Proceedings of the National Academy of Sciences, 109, pp. 14726-14727 (September 11, 2012).
- 10. Schär et al., "The role of increasing temperature variability in European summer heatwaves," Nature 427, pp. 332-336 (2004).
- 11. Hansen et al., "Global surface temperature change," Review of Geophysics 48, RG4004 (2010).
- 12. Steig et al., "Warming of the Antarctic ice-sheet surface since the 1957 International Geophysical Year," Nature 457, 459-462 (22 January 2009); Orsi et al., Little Ice Age cold interval in West Antarctica: Evidence from borehole temperature at the West Antarctic Ice Sheet (WAIS) Divide, Geophysical Research Letters 39, L09710 (2012); Muto et al., "Recent surface temperature trends in the interior of East Antarctica from borehole firn temperature measurements and geophysical inverse methods," Geophysical Research Letters 38, L15502 (2011).
- 13. Smith et al., "Improved surface temperature prediction for the coming decade from a global climate model," Science 317 (5839), pp. 796–799 (2007); Keenlyside et al., "Advancing decadal-scale climate prediction in the North Atlantic sector," Nature 453, pp. 84-88 (2008).
- 14. Gleckler et al., "Human-induced global ocean warming on multidecadal timescales," Nature Climate Change 2, pp. 524–529 (2012); Jones and Stott, "Sensitivity of the attribution of near surface temperature warming to the choice of observational dataset," Geophysical Research Letters 38, L21702 (2011); Huber and Knutti, "Anthropogenic and

natural warming inferred from changes in Earth's energy balance," Nature Geoscience 5, pp. 31-36 (2012).

- 15. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 10: "Detection and Attribution of Climate Change: from Global to Regional," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b FinalDraft Chapter10.pdf.
- 16. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 2: "Observations: Atmosphere and Surface," available at www.climatechange2013.org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_ Chapter02.pdf.
- 17. Meehl et al., "Externally Forced and Internally Generated Decadal Climate Variability Associated with the Interdecadal Pacific Oscillation," Journal of Climate 26, pp. 7298-7310 (2013).
- 18. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, "Summary for Policymakers," available at www.climatechange2013. org/images/uploads/WGI_AR5_SPM_brochure.pdf; Kaufmann, et al., "Reconciling anthropogenic climate change with observed temperature 1998-2008," Proceedings of the National Academy of Sciences of the United States of America 108 (29), pp. 11790-11793, 19 July 2011. www.pnas.org/content/108/29/11790.full.
- Meehl et al., "Externally Forced and Internally Generated Decadal Climate Variability 19. Associated with the Interdecadal Pacific Oscillation," Journal of Climate 26, pp. 7298-7310 (2013): Guemas et al., "Retrospective prediction of the global warming slowdown in the past decade," Nature Climate Change 3, pp. 649-653 (2013); Kosaka and Xie, "Recent global-warming hiatus tied to equatorial Pacific surface cooling," Nature 501, pp. 403-407 (2013)
- 20. Kosaka and Xie, "Recent global-warming hiatus tied to equatorial Pacific surface cooling," Nature 501, pp. 403-407 (2013).
- Meehl et al., "Model-based evidence of deep-ocean heat uptake during surface-21. temperature hiatus periods," Nature Climate Change 1, 360-364 (2011).
- 22. Rosenfeld et al., "Aerosol Cloud-Mediated Radiative Forcing: Highly Uncertain and Opposite Effects from Shallow and Deep Clouds," Climate Science for Serving Society, pp. 105-149 (2003).
- 23. Stan et al., "An ocean-atmosphere climate simulation with an embedded cloud resolving model," Geophysical Research Letters 37, L01702 (2010).
- 24. Lauer et al., "The impact of global warming on marine boundary layer clouds over the Eastern Pacific—A regional model study," Journal of Climate 23, pp. 5844–5863 (2010); Clement et al., "Observational and model evidence for positive low-level cloud feedback," Science 325, pp. 460–464 (2009); Dessler, "A Determination of the Cloud Feedback from Climate Variations over the Past Decade," Science 330, pp. 1523–1527 (2010).
- 25. Swanson et al., "Long-term natural variability and 20th century climate change," Proceedings of the National Academy of Sciences (2009); DelSole et al., "A significant component of unforced multidecadal variability in the recent acceleration of global warming," Journal of Climate 24 (3), pp. 909-926 (2011); Latif and Keenlyside, "A perspective on decadal climate variability and predictability," Deep-Sea Research II 58, pp. 1880-1894 (2011).
- 26. Wu et al., "On the time-varying trend in global-mean surface temperature," *Climate* Dynamics 37, pp. 759-773 (2011).
- 27. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 14: "Climate Phenomena and their Relevance for Future Regional Climate Change," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b_FinalDraft_Chapter14.pdf.
- 28 This cycle is driven by the sun's internal magnetic dynamo: as magnetic activity increases, it opens up hot areas called faculae, which make the sun slightly brighter.
- 29. Foukal et al., "Variations in solar luminosity and their effect on the Earth's climate," Nature 443, p 161 (2006).
- 30. Lean et al., "Reconstructions of solar irradiance since 1610: Implications for climate change," Geophysical Research Letters, vol. 22, pp. 3195-3198 (1995).
- Wang et al., "Modeling the sun's magnetic field and irradiance since 1713," The 31. Astrophysical Journal, vol. 625, pp. 522-538 (2005).

- 32. Douglass and Clader, "Climate sensitivity of the Earth to solar irradiance," Geophysical Research Letters, vol. 29, pp. 33-1 to 33-4 (2002).
- 33. Frölich and Lean, "Solar radiative output and its variability: evidence and mechanisms," The Astronomy and Astrophysics Review, vol. 12, pp. 273-320 (2004); Lean, "Cycles and trends in solar irradiance and climate," Wiley Interdisciplinary Reviews: Climate Change 1, pp. 111–122 (2010).
- Haigh et al., "An influence of solar spectral variations on radiative forcing of climate," 34. Nature 467, pp. 696-699 (2010).
- CLOUD collaboration, "2009 Progress report on PS215/CLOUD," (2010). Available at 35. http://cds.cern.ch/record/1257940/files/SPSC-SR-061.pdf.
- 36. One of the most extreme warming events known in the history of Earth was the Paleocene-Eocene Thermal Maximum, which took place about 56 million years ago. The evidence is that this warming was accompanied by a massive release of CO₂, but recent research suggests that it took place slower than the current increase in CO₂. Cui et al., "Slow release of fossil carbon during the Palaeocene-Eocene Thermal Maximum," Nature Geoscience 4, pp. 481-485 (2011).
- 37. Barnola et al., "Historical Carbon Dioxide Record from the Vostok Ice Core," available at http://cdiac.ornl.gov/trends/co2/vostok.html.
- 38. For example see Jones, G. S. and P. A. Stott (2011), "Sensitivity of the attribution of near surface temperature warming to the choice of observational dataset", Geophysical Research Letters, 38, L2170; Huber and Knutti, "Anthropogenic and natural warming inferred from changes in the Earth's energy balance," Nature Geoscience 5, 31–36 (2012); Foster and Rahmstorf, "Global temperature evolution, 1979–2010," Environmental Research Letters 6, 044022 (2011); Gleckler et al., "Human-induced global ocean warming on multidecadal timescales," Nature Climate Change 2, pp. 524-529 (2012).
- 39. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 10: "Detection and Attribution of Climate Change: from Global to Regional," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b FinalDraft Chapter10.pdf.
- 40. Bond et al., "Bounding the role of black carbon in the climate system: A scientific assessment," Journal of Geophysical Research: Atmospheres 118, pp. 5380-5582 (2013).
- "Towards new scenarios for analysis of emissions, climate change, impacts, and response 41. strategies," IPCC Expert Meeting Report (19-21 September 2007, Noordwijkerhout, The Netherlands). IIASA, RCP Database (Version 2.0), available at www.iiasa.ac.at/web-apps/ tnt/RcpDb/dsd?Action=htmlpage&page=welcome; van Vuuren, et al., "The representative concentration pathways: an overview," Climatic Change 109, pp. 5-31 (2011).
- 42. Moss et al., "The next generation of scenarios for climate change research and assessment," Nature 462 (2010); van Vuuren, et al., "The representative concentration pathways: an overview," Climatic Change 109, pp. 5–31 (2011).
- IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical 43. Science Basis, "Summary for Policymakers," available at www.climatechange2013. org/images/uploads/WGI AR5 SPM brochure.pdf; Allen et al., "Warming caused by cumulative carbon emissions towards the trillionth tonne," Nature (2009); Meinshausen et al., "Greenhouse-gas emission targets for limiting global warming to 2°C," Nature 458, 1158-1162 (2009).
- 44. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 2: "Observations: Atmosphere and Surface," available at www.climatechange2013.org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_ Chapter02.pdf.
- 45. Meinshausen et al., "The RCP greenhouse gas concentrations and their extensions from 1765 to 2300," Climatic change 109, pp. 213–241, (2011).
- Mann, "Defining dangerous anthropogenic interference," PNAS 106, pp. 4065–4066 46 (10 March 2009)
- 47. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 12: "Long-term Climate Change: Projections, Commitments and Irreversibility," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b_FinalDraft_Chapter12.pdf.
- Smith, et al., "Assessing dangerous climate change through an update of the 48. Intergovernmental Panel on Climate Change (IPCC) 'reasons for concern,'" Proceedings of the National Academy of Sciences 106, pp. 4133–4137 (2009).

- 49. Collins et al., "Development and evaluation of and Earth-system model – HadGEM2," Geoscientific Model Development Discussions 4, pp. 997–1062 (2011).
- 50. Masson and Knutti, "Climate model genealogy," Geophysical Research Letters 38, p L08703 (2011).
- 51. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, "Summary for Policymakers," available at www.climatechange2013.org/ images/uploads/WGI_AR5_SPM_brochure.pdf.
- 52. Gordon et al., "An Observationally Based Constraint on the Water-Vapor Feedback," Journal of Geophysical Research: Atmospheres, published online: http://onlinelibrary. wiley.com/doi/10.1002/2013JD020184/abstract
- 53. Global warming is sometimes encapsulated in a single number, "climate sensitivity", although this can be defined in different ways. Equilibrium climate sensitivity is the long-term temperature change produced by a doubling of atmospheric CO₂, and recent estimates range from about 1.5°C to 4.5°C. Equilibrium climate sensitivity does not reflect short-temperature change, because of the heat-buffering effect of the oceans – so arguably it is not very relevant for policy. Transient climate sensitivity is the temperature rise when the atmospheric CO₂ concentration has doubled in a scenario where concentration increases at 1% per year. According to the AR5, transient climate sensitivity is likely to be between 1.0 and 2.5°C - although one recent study (Otto et al., below) found that the transient climate sensitivity could be slightly lower than most previous estimates, with a range from 0.9 to 2°C. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 12: "Long-term Climate Change: Projections, Commitments and Irreversibility," available at www.climatechange2013. org/images/uploads/WGIAR5 WGI-12Doc2b FinalDraft Chapter12.pdf; Schmittner et al., "Climate sensitivity estimated from temperature reconstructions of the last glacial maximum," Science 334, pp. 1385-1388 (9 December 2011); Köhler et al., "What caused Earth's temperature variations during the last 800,000 years? Data-based evidence on radiative forcing and constraints on climate sensitivity," Quaternary Science Reviews 29 (1-2), pp. 129–145 (2010); Otto et al.,"Energy budget constraints on climate response," Nature Geoscience 6, pp. 415-416 (2013).
- 54. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 4: "Observations: Cryosphere," available at www.climatechange2013.org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_ Chapter04.pdf.
- 55. Van den Broeke et al., "Partitioning Recent Greenland Mass Loss," Science 326, pp. 984-986
- For example, Laxon et al., "CryoSat-2 estimates of Arctic sea ice thickness and volume," 56. Geophysical Research Letters 40, pp. 732-737 (2013).
- 57. Kinnard et al., "Reconstructed changes in Arctic sea ice over the past 1,450 years," Nature 479, pp. 509–512 (2011)
- IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical 58. Science Basis, Chapter 10: "Detection and Attribution of Climate Change: from Global to Regional," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b_FinalDraft_Chapter10.pdf.
- 59 Pritchard et al., "Antarctic ice-sheet loss driven by basal melting of ice shelves," Nature 484, pp. 502-505 (2012).
- 60. Warneke et al., "An important contribution to springtime Arctic aerosol from biomass burning in Russia," Geophysical Research Letters 37, L01801 (2010).
- 61. Hadley and Kirchstetter, "Black-carbon reduction of snow albedo," Nature Climate Change 2, pp. 437-440 (2012).
- 62. Picard et al., "Inhibition of the positive snow-albedo feedback by precipitation in interior Antarctica," Nature Climate Change 2, pp. 795–798 (2012).
- 63. Shepherd et al., "A Reconciled Estimate of Ice-sheet Mass Balance," Science 338, pp. 1183-1189 (2012).
- 64. IPCC, Fourth Assessment Report: Climate Change 2007, Working Group II: Impacts, Adaptation and Vulnerability, "Fresh water resources and their management," available at http://ipcc.ch/publications_and_data/ar4/wg2/en/ch3s3-es.html
- "Snow, Water, Ice, Permafrost in the Arctic: Climate Change and the Cryosphere," 65. available at www.amap.no/swipa.

- 66. Screen and Simmonds, "The central role of diminishing sea ice in recent Arctic temperature amplification," Nature 464, pp. 1337–1337 (29 April 2010); Hansen et al., "Global surface temperature change," Review of Geophysics 48, RG4004 (2010); Screen et al., "Local and remote controls on observed Arctic warming," Geophysical Research Letters 39, L10709 (2012).
- 67. DeConto et al., "Past extreme warming events linked to massive carbon release from thawing permafrost," Nature 484 (7392), pp. 87-91 (2012); Schuur et al., "Vulnerability of permafrost carbon to climate change: implications for the global carbon cycle," BioScience 58(8), pp. 701-714 (2008); Schuur et al., "The effect of permafrost thaw on old carbon release and net carbon exchange from tundra," Nature 459, pp. 556–559 (2009).
- 68. Wang and Overland, "A sea ice free summer Arctic within 30 years: An update from CMIP5 models." Geophysical Research Letters 39. L18501 (2012).
- 69. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 12: "Long-term Climate Change: Projections, Commitments and Irreversibility," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b FinalDraft Chapter12.pdf.
- 70. Shakhova et al., "Extensive methane venting to the atmosphere from sediments of the East Siberian Arctic Shelf," Science 327, pp. 1246–1250 (2010). Anthony, et al., "Geologic methane seeps along boundaries of Arctic permafrost thaw and melting glaciers," Nature Geoscience 5, pp. 419–426 (published online 20 May 2012)
- 71. Holland and Kwok, "Wind-driven trends in Antarctic sea-ice drift," Nature Geoscience 5, pp. 872-875 (2012).
- 72. Pritchard et al., "Antarctic ice-sheet loss driven by basal melting of ice shelves," Nature 484, pp. 502-505 (2012).
- 73. See for example Rignot et al., "Ice flow of the Antarctic Ice Sheet," Science 333, pp. 1427-1430 (9 September 2011).
- 74. The following studies measure sea level rise relative to slightly different reference points - for example mean sea level between 1980 and 2000, or between 1986 and 2005 - but the differences are small, and for simplicity they can all be considered as rises above sea level in 1990. Stefan Rahmstorf, "A Semi-Empirical Approach to Projecting Future Sea-Level Rise," Science 315, pp. 368–370, 19 January 2007; Martin Vermeer and Stefan Rahmstorf, "Global sea level linked to global temperature," Proceedings of the National Academy of Sciences 106, pp. 21527–21532 (2009). Pfeffer et al., "Kinematic constraints on glacier contributions to 21st-century sea-level rise," Science 321, pp. 1340–1343 (2008); Jevrejeva et al., "Sea level projections to AD2500 with a new generation of climate change scenarios, Global and Planetary Change 80, pp. 14–20 (2012).; Schaeffer et al., "Long-term sea-level rise implied by 1.5°C and 2°C warming levels," Nature Climate Change 2, pp. 867–870 (2012); Meehl et al., "Relative outcomes of climate change mitigation related to global temperature versus sea-level rise," Nature Climate Change 2, pp. 576–580 (2012); Moore et al., "Semi-empirical and process-based sea level projections," Reviews of Geophysics 51, pp. 484–522 (2013).
- 75. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 13: "Sea level change," available at www.climatechange2013.org/ images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_Chapter13.pdf.
- 76. Moore et al., "Semi-empirical and process-based sea level projections," Reviews of Geophysics 51, pp. 484-522 (2013).
- 77. Meehl et al., "Relative outcomes of climate change mitigation related to global temperature versus sea-level rise," Nature Climate Change 2, pp. 576–580 (2012); Schaeffer et al., "Long-term sea-level rise implied by 1.5°C and 2°C warming levels," Nature Climate Change 2, pp. 867-870 (2012).
- Intergovernmental Panel on Climate Change Workshop on Sea Level Rise and Ice Sheet 78. Instabilities [Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S. Allen, and P.M. Midgley (eds.)], available at www.ipcc.ch/pdf/supporting-material/SLW WorkshopReport kuala lumpur.pdf
- Jevrejeva et al., "Sea level projections to AD2500 with a new generation of climate 79. change scenarios, Global and Planetary Change 80, pp. 14-20 (2012).
- 80. The far larger East Antarctic ice sheet is much more stable, and is not expected to collapse even millennia from now. IPCC, Fourth Assessment Report: Climate Change 2007, Working Group I: The Physical Science Basis, "Observations: Changes in Snow, Ice and Frozen Ground," available at http://ipcc.ch/publications_and_data/ar4/wg1/en/ch4s4-1.html.

- IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, "Summary for Policymakers," available at www.climatechange2013.org/ images/uploads/WGI_AR5_SPM_brochure.pdf.
- IPCC, Fourth Assessment Report: Climate Change 2007, Synthesis Report, "Summary for Policymakers," available at www.climatechange2013.org/images/uploads/WGI_AR5_ SPM_brochure.pdf.
- Hallegatte at al., "Future flood losses in major coastal cities," Nature Climate Change 3, pp. 802–806 (2013).
- 84. Sallenger et al., "Hotspot of accelerated sea-level rise on the Atlantic coast of North America," *Nature Climate Change* 2, pp 884–888 (2012).
- 85. Wada et al., "Global depletion of groundwater resources," *Geophysical Research Letters* 37, L20402 (2010); Konikow, "Contribution of global groundwater depletion since 1900 to sea-level rise," *Geophysical Research Letters* 38, L17401 (2011); Church et al., "Revisiting the Earth's sea-level and energy budgets from 1961 to 2008," *Geophysical Research Letters* 38, L18601 (2011).
- 86. Rahmstorf, "What makes sea-level rise?" RealClimate (1 June 2012), available at www.realclimate.org/index.php/archives/2012/06/what-makes-sea-level-rise; Wada et al., "Past and future contribution of global groundwater depletion to sea-level rise," *Geophysical Research Letters* 39, L09402 (2012); Rahmstorf et al., "Testing the robustness of semi-empirical sea level projections," *Climate Dynamics* (2011).
- IPCC, Managing the Risks of Extreme Events and Disasters to Advance Climate Change Adaptation, eds. Field, et al., Cambridge University Press, pp. 111–113, p. 119 (2012).
- 88. Alexander, et al., "Global observed changes in daily climate extremes of temperature and precipitation," *Journal of Geophysical Research* 111, D05109 (2006).
- IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 12: "Long-term Climate Change: Projections, Commitments and Irreversibility," available at www.climatechange2013.org/images/uploads/WGIAR5_ WGI-12Doc2b_FinalDraft_Chapter12.pdf.
- Dai, "Characteristics and trends in various forms of the Palmer Drought Severity Index during 1900–2008," *Journal of Geophysical Research* 116, D12115 (2011); IPCC, *Managing the Risks of Extreme Events and Disasters to Advance Climate Change Adaptation*, eds. Field, et al., Cambridge University Press, pp. 111–113, p. 119 (2012).
- Liu and Allan, "Observed and simulated precipitation responses in wet and dry regions, 1850–2100," *Environmental Research Letters* 8 p 34002 (2013). Hirabayashi et al., "Global flood risk under climate change," *Nature Climate Change* 3, pp. 816–821 (2013).
- Sherwood and Huber, "An adaptability limit to climate change due to heat stress," Proceedings of the National Academy of Sciences 107, pp. 9552–9555 (2010).
- 93. Gerten at al., "Asynchronous exposure to global warming: freshwater resources and terrestrial ecosystems," *Environmental Research Letters* 8, p 34032 (2013).
- 94. Battisti and Naylor, "Historical Warnings of Future Food Insecurity with Unprecedented Seasonal Heat," *Science* 323, pp. 240–244 (2009).
- 95. Stott, et al., "Human contribution to the European heatwave of 2003," *Nature* 432, pp. 610–614 (2004); Otto, et al., "Reconciling two approaches to the attribution of the 2010 Russian heat wave," *Geophysical Research Letters* 39, L04702 (2012); Min, et al., "Human contribution to more-intense precipitation extremes," *Nature* 470, pp. 378–381 (2011); Pall, et al., "Anthropogenic greenhouse gas contribution to flood risk in England and Wales in autumn 2000, *Nature* 470, pp. 382–386 (2011); Coumou and Rahmstorf, "A decade of weather extremes," *Nature Climate Change* 2, pp.491–496 (2012).
- 96. IPCC, Managing the Risks of Extreme Events and Disasters to Advance Climate Change Adaptation, eds. Field, et al., Cambridge University Press, pp. 111–113, p. 119 (2012); Alexander, et al., "Global observed changes in daily climate extremes of temperature and precipitation," Journal of Geophysical Research 111, D05109 (2006).
- Liu and Allan, "Observed and simulated precipitation responses in wet and dry regions, 1850–2100," *Environmental Research Letters* 8 p 34002 (2013). Hirabayashi et al., "Global flood risk under climate change," *Nature Climate Change* 3, pp. 816–821 (2013).
- Min et al., "Human contribution to more-intense precipitation extremes," Nature 470, pp. 378–381 (2011).
- Holland and Bruyere, "Recent intense hurricane response to global climate change," Climate Dynamics, published online, DOI: 10.1007/s00382-013-1713-0 (2013) http://link. springer.com/article/10.1007/s00382-013-1713-0].

- 100. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 13: "Sea level change," available at www.climatechange2013.org/ images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_Chapter13.pdf.
- 101. IPCC, Managing the Risks of Extreme Events and Disasters to Advance Climate Change Adaptation, eds. Field, et al., Cambridge University Press (2012), p. 269.
- 102. Hirabayashi et al., "Global flood risk under climate change," *Nature Climate Change* 3, pp. 816–821 (2013).
- Alves, "Numerical modeling of ocean swell contributions to the global wind-wave climate," Ocean Modelling 11, 98–122 (2006).
- 104. Bitner-Gregerson et al., "Ship and Offshore Structure Design in Climate Change Perspective," *Springer*, ISBN 978-3-642-34137-3 (2013).
- 105. Toffioli et al., "Extreme Waves in Random Crossing Seas: Laboratory Experiments and Numerical Simulations," *Geophysical Research Letters* 38, L06605 (2011).
- 106. For example Wang et al., "North Atlantic wave height trends as reconstructed from the Twentieth Century Reanalysis," *Geophysical Research Letters* 39, p. L18705 (2012); Young et al., "Global trends in wind speed and wave height," *Science* 332, pp. 451–455 (2011).
- Bitner-Gregerson et al., "Ship and Offshore Structure Design in Climate Change Perspective," Springer, ISBN 978-3-642-34137-3 (2013).
- 108. Bitner-Gregerson et al., "Potential Impact of Climate Change on Tanker Design," presentation at the OMAE 2011 Conference, Rotterdam June 2011.
- 109. Byrnes et al., "Climate-driven increases in storm frequency simplify kelp forest food webs," *Global Change Biology* 17, pp. 2513–2524 (2011).
- 110. Le Quéré et al., "Impact of climate change and variability on the global oceanic sink of CO₂," *Global Biogeochemical Cycles* 24, GB4007 (2010).
- 111. Scientists mark acidity on the pH scale, a measure of the concentration of hydrogen ions in a solution. It is an inverse logarithmic scale, meaning that a solution with a pH of 5, for example, has 10 times the concentration of hydrogen ions as a solution with a pH of 6. The lower the number, the more acid the solution. The sea's pre-industrial pH was about 8.25 (which is less acid, or more alkaline, than neutral distilled water with a pH of 7). The sea's pH fell to 8.13 by the 1990s – an increase of more than 30% in the concentration of hydrogen ions.
- 112. Bernie et al., "Influence of mitigation policy on ocean acidification," *Geophysical Research Letters* 7, p. L15704 (2010).
- 113. Bernie et al., "Influence of mitigation policy on ocean acidification," *Geophysical Research Letters* 7, p. L15704 (2010); Caldeira and Wickett, "Oceanography: Anthropogenic carbon and ocean pH," *Nature* 425, p. 365 (2003); Raven et al., "Ocean acidification due to increasing atmospheric carbon dioxide," *The Royal Society*, Policy document 12/05 (2005); Turley et al., "Reviewing the Impact of Increased Atmospheric CO₂ on Oceanic pH and the Marine Ecosystem," in *Avoiding Dangerous Climate Change*, pp. 65–70, Cambridge University Press (2006); Hönisch et al., "The geological record of ocean acidification," *Science* 335, pp. 1058–1063 (2012).
- 114. Riebesell, "Sensitivities of marine carbon fluxes to ocean change," *Proceedings of the National Academy Sciences* 106, pp. 20602–20609 (2009).
- 115. Six et al,. "Global warming amplified by reduced sulphur fluxes as a result of ocean acidification," *Nature Climate Change* 3, pp. 975–978 (2013).
- 116. Flynn et al., "Changes in pH at the exterior surface of plankton with ocean acidification," *Nature Climate Change* 2, pp. 510–513 (2012).
- 117. Feely et al., "Evidence for upwelling of corrosive 'acidified' water onto the continental shelf," *Science* 320, pp. 1490–1492 (2008).
- 118. Available at www.stateoftheocean.org/pdfs/IPSO-Summary-Oct13-FINAL.pdf.
- 119. Steinacher, et al., "Imminent ocean acidification in the Arctic projected with the NCAR global coupled carbon cycle-climate model," *Biogeosciences* 6, pp. 515–533 (2009); McNeil and Matear, "Southern Ocean acidification: A tipping point at 450-ppm atmospheric CO₂," *Proceedings of the National Academy of Sciences* 105, pp. 18860–18864, 20 November 2008; Yamamoto-Kawai et al., "Aragonite undersaturation in the Arctic Ocean: Effects of ocean acidification and sea ice melt," *Science* 326, 1098–1100 (2009).
- 120. Hönisch et al., "The geological record of ocean acidification," *Science* 335, pp. 1058–1063 (2012).

- 121. Six et al., "Global warming amplified by reduced sulphur fluxes as a result of ocean acidification," Nature Climate Change 3, pp. 975–978 (2013).
- 122. UN-Water, Managing Water under Uncertainty and Risk, The United Nations World Water Development Report 4 (2012); Milly et al., "Global pattern of trends in streamflow and water availability in a changing climate," Nature 438, pp. 347–350 (2005); Vorosmarty et al., "Global water resources: vulnerability from climate change and population growth," Science 289, pp. 284-2888 (2000).
- 123. IPCC, Fourth Assessment Report: Climate Change 2007, Working Group II: Impacts, Adaptation and Vulnerability, "Fresh water resources and their management," available at http://ipcc.ch/publications and data/ar4/wg2/en/ch3s3-es.html; Liu and Allan, "Observed and simulated precipitation responses in wet and dry regions, 1850–2100," Environmental Research Letters 8, pp. 34002 (2013).
- 124. Dai, "Drought under global warming: a review," Wiley Interdisciplinary Reviews: Climate Change, DOI: 10.1002/wcc.81 (2011).
- 125. Immerzee et al., "Climate Change Will Affect the Asian Water Towers," Science 328, pp. 1382-1385 (2010).
- 126. Barnett et al., "Potential impacts of a warming climate on water availability in snowdominated regions," Nature 438, pp. 303–309 (2005); Fraser, "Melting in the Andes: Goodbye glaciers," Nature 491, pp. 180–182 (2012).
- 127. For example, Piao et al. "The impacts of climate change on water resources and agriculture in China," Nature Climate Change 467, pp. 43–51 (2010).
- 128. Lobell, et al., "Climate trends and global crop production since 1980," Science 333, pp. 616-620 (29 July 2011); Lobell et al, "Extreme heat effects of wheat senescence in India," Nature Climate Change 2, pp. 186-189 (2012); Welch et al., "Rice yields in tropical/ subtropical Asia exhibit large but opposing sensitivities to minimum and maximum temperatures," Proceedings of the National Academy of Sciences 107, pp. 107: 14562-14567 (2010); Battisti and Naylor, "Historical Warnings of Future Food Insecurity with Unprecedented Seasonal Heat," Science 323, pp. 240–244 (2009).
- 129. Battisti and Naylor, "Historical Warnings of Future Food Insecurity with Unprecedented Seasonal Heat," Science 323, pp. 240-244 (2009)
- 130. IPCC, Fourth Assessment Report: Climate Change 2007, Working Group II: Impacts, Adaptation and Vulnerability, "Food, fibre and forest products," available at http://ipcc. ch/publications_and_data/ar4/wg2/en/ch5s5-es.html.
- 131. IPCC, Fourth Assessment Report: Climate Change 2007, Synthesis Report, available at http://ipcc.ch/publications_and_data/ar4/syr/en/figure-3-6.html.
- 132. Lobell et al., "Climate trends and global crop production since 1980," Science 333, pp. 616–620 (2011).
- 133. Patz et al., "Impact of regional climate change on human health," Nature 438, pp. 310-317 (2005).
- 134. Baumann et al., "Reduced early life growth and survival in a fish in direct response to increased carbon dioxide," Nature Climate Change 2, pp. 38-41 (2012); Frommel et al., "Severe tissue damage in Atlantic cod larvae under increasing ocean acidification," Nature Climate Change 2, pp. 42–46 (2012); Jacobsen, et al., "Biodiversity under threat in glacierfed river systems," Nature Climate Change 2, pp. 361-364 (2012).
- 135. Wassmann, "Footprints of climate change in the Arctic marine ecosystem," Global Change Biology 17, pp. 1235-1249 (2011).
- 136. Van Vliet et al., "Vulnerability of US and European electricity supply to climate change," Nature Climate Change 2, pp. 676-681 (2012).
- 137. See for example Madani and Lund, "Estimated impacts of climate warming on California's high-elevation hydropower," Climatic Change 102, pp. 521-538 (2010).
- 138. Hertel et al., "The poverty implications of climate-induced crop yield changes by 2030," Global Environmental Change, GTAP Working Paper 59 (2010); Nelson et al., Food Security, Farming, and Climate Change to 2050: scenarios, results, policy options, International Food Policy Research Institute (2010); Willenbockel, Exploring Food Price Scenarios Towards 2030 with a Global Multi-Region Model, Oxfam Research Report (1 June 2011), available at www.oxfam.org/sites/www.oxfam.org/files/rr-exploring-foodprice-scenarios-010611-en.pdf.
- 139. Lobell et al., "Climate trends and global crop production since 1980," Science 333, pp. 616-620 (2011).

- 140. Alley et al., "Abrupt Climate Change," Science 299, pp. 2005–2010 (2003).; Alley et al., "Abrupt Climate Change – Inevitable Surprises," Committee on abrupt climate change, National Research Council (2002). National Academy Press, Washington, D.C. 230pp.
- 141. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 12: "Long-term Climate Change: Projections, Commitments and Irreversibility," available at www.climatechange2013.org/images/uploads/WGIAR5 WGI-12Doc2b_FinalDraft_Chapter12.pdf.
- 142. Lenton et al., "Tipping elements in the Earth's climate system," Proceedings of the National Academy of Sciences 105, pp. 1786-1793 (2008).
- 143. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 1: "Introduction," available at www.climatechange2013.org/ images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_Chapter01.pdf.
- 144. Lenton, "Early warning of climate tipping points," Nature Climate Change 1, pp. 201–209 (2011); Lenton, "Future climate surprises," in The Future of the World's Climate, ed. Henderson-Sellers and McGuffie, pp. 489–507 (2012).
- 145. Kriegler et al., "Imprecise probability assessment of tipping points in the climate system," Proceedings of the National Academy of Sciences 106, pp. 5041–5046 (2009).
- 146. Pfeffer et al., "Kinematic constraints on glacier contributions to 21st century sea-level rise," Science 321, pp. 1340–1343 (2008); Bamber et al., "Reassessment of the potential sea-level rise from a collapse of the West Antarctic Ice Sheet," Science 324, pp. 901-903 (2009).
- 147. Cook and Vizy, "Effects of twenty-first-century climate change on the Amazon Rain Forest," Journal of Climate 21, pp. 542-560 (2008).
- 148. Lenton et al., "Tipping elements in the Earth's climate system," Proceedings of the National Academy of Sciences 105, pp. 1786–1793 (2008).
- 149. Turner and Annamalai, "Climate change and the South Asian summer monsoon," Nature Climate Change 2, pp. 587-595 (2012).
- 150. Bollasina et al., "Anthropogenic aerosols and the weakening of the South Asian Summer Monsoon," Science 334, pp. 502-505 (2011); Ramanathan et al., "Atmospheric brown clouds: impacts on South Asian climate and hydrological cycle," Proceedings of the National Academy of Sciences 102, pp. 5326-5333 (2005).
- 151. Allison et al., The Copenhagen Diagnosis, 2009: Updating the world on the latest climate science, The University of New South Wales Climate Change Research Centre (CCRC), Sydney, Australia, p. 41. Available at www.copenhagendiagnosis.com.
- 152. Lenton, "Future climate surprises," in The Future of the World's Climate, ed. Henderson-Sellers and McGuffie, p. 494 (2012).
- 153. Robinson et al., "Multistability and critical thresholds of the Greenland ice sheet," Nature Climate Change 2, pp. 429-432 (2012).
- 154. Box et al., "Greenland ice sheet albedo feedback: thermodynamics and atmospheric drivers," in press; NASA, "Satellites see unprecedented Greenland ice sheet surface melt" (24 July 2012), available at www.nasa.gov/topics/earth/features/greenland-melt.html.
- 155. Henderson-Sellers, ed., The Future of the World's Climate: A Modelling Perspective (2012), pp. 494–495; Zimov et al., "Permafrost and the global carbon budget," Science 312, pp. 1612-1613 (2006)
- 156. Schuur and Abbott, "High risk of permafrost thaw," Nature 480, pp. 32–33 (2011).
- 157. Khvorostyanov et al., "Vulnerability of east Siberia's frozen carbon stores to global warming," Geophysical Research Letters 35, L10703 (2008).
- 158. Shakhova et al., "Extensive methane venting to the atmosphere from sediments of the East Siberian Arctic Shelf," Science 327, pp. 1246-1250 (2010).
- 159. Anthony et al., "Geologic methane seeps along boundaries of Arctic permafrost thaw and melting glaciers," Nature Geoscience 5, pp. 419-426 (2012).
- 160. Pires and Costa, "Deforestation causes different subregional effects on the Amazon bioclimatic equilibrium," Geophysical Research Letters 40, pp. 3618-3623 (2013).
- 161. Pearce, "Arctic sea ice may have passed crucial tipping point," New Scientist 27 March 2012.
- 162. www.ipcc.ch/report/ar5/wg1
- 163. Another factor that has kept this range of projections wide is that more elements of the climate system are being included in models, so the uncertainties associated with each component are included explicitly in the model outputs.

- 164. www.argo.ucsd.edu
- 165. Myhre et al., "Radiative forcing of the direct aerosol effect from AeroCom Phase II simulations," *Atmospheric Chemistry and Physics* 13, pp. 1853–1877 (2013).
- 166. The two most widely used global average temperature records are published by the US National Aeronautics and Space Administration (NASA), http://data.giss.nasa.gov/gistemp/, and the University of East Anglia's Climate Research Unit, www.cru.uea.ac.uk/cru/info/warming/.
- 167. National Snow & Ice Data Center's sea ice index, http://nsidc.org/data/seaice_index/.
- Mann et al., "Proxy-based reconstructions of hemispheric and global surface temperature variations over the past two millennia," *Proceedings of the National Academy of Sciences* 105, pp. 13252–13257 (2008); Jones, et al., "The evolution of climate over the last millennium," *Science* 292, pp. 662–667 (2001).
- 169. Caillon, et al., "Timing of atmospheric CO₂ and Antarctic temperature changes across Termination III," *Science* 299, pp. 1728–1731 (2003).
- 170. Severinghaus, "What does the lag of CO₂ behind temperature in ice cores tell us about global warming?" originally posted 3 December 2004, www.realclimate.org/index.php/archives/2004/12/co2-in-ice-cores/.
- 171. Shakun et al., "Global warming preceded by increasing carbon dioxide concentrations during the last deglaciation," *Nature* 484, pp. 49–54 (2012).
- 172. IPCC, Fifth Assessment Report: Climate Change 2013, Working Group I: The Physical Science Basis, Chapter 6: "Carbon and other Biogeochemical Cycles," available at www.climatechange2013.org/images/uploads/WGIAR5_WGI-12Doc2b_FinalDraft_ Chapter06.pdf.
- 173. International Energy Agency, "Global carbon-dioxide emissions increase by 1.0 Gt in 2011 to record high," 24 May 2012, www.iea.org/newsroomandevents/news/2012/may/ name,27216,en.html.
- 174. Steig et al., "Warming of the Antarctic ice-sheet surface since the 1957 International Geophysical Year," *Nature* 457, 459–462 (2009).
- 175. Russell et al., *The Independent Climate Change E-mails Review* (July 2010) available at www.cce-review.org/pdf/FINAL%20REPORT.pdf; UK House of Commons Science and Technology Select Committee, *The Reviews into the University of East Anglia's Climatic Research Unit's E-mails* (January 2011) available at www.publications.parliament.uk/pa/cm201011/cmselect/cmsctech/444/444.pdf; Scientific Assessment Panel, *Report of the International Panel set up by the University of East Anglia to examine the research of the Climatic Research Unit* (April 2010) available at www.uea.ac.uk/mac/comm/media/press/crustatements/sap; US Environmental Protection Agency, *Denial of Petitions for Reconsideration of the Endangerment and Cause or Contribute Findings for Greenhouse Gases under Section 202(a) of the Clean Air Act* (July 2010) available at http://epa.gov/climatechange/endangerment/petitions.html.